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## MINUTES OF THE MEETING STEERING COMMITTEE (SC)

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Meeting No<sup>o</sup> 70

**Wednesday, January 29, 2025**

9:00 AM to 12:00 PM

In Person – Maison du Développement Durable- 50 rue Sainte Catherine  
Montreal

And Videoconference – TEAMS

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<b>Present:</b>	Felix Boulanger	Hydro-Québec
	Marc Dunn	Niskamoon Corporation
	Luc Duquette	Hydro-Québec
	Louie Kanatewat	Cree Nation of Chisasibi
	Mélanie Leblanc	Niskamoon Corporation
	Josée Lefebvre	Canadian Wildlife Service
	Marie-Eve Lemieux	Hydro-Québec
	Geraldine Mark	Cree Nation of Wemindji
	Ernest Moses	Cree Nation of Waskaganish
	Mimie Neacappo	Niskamoon Corporation
	John Lameboy	Cree Nation of Chisasibi
	Manon Sorais	Eeyou marine region wildlife board
	Ernie Rabbitskin	Niskamoon Corporation
	Robbie Tapiatic	Cree Nation of Chisasibi
Cassandra Weapenicappo	Cree Nation of Eastmain	
<b>Guest:</b>	Paul del Giorgio	UQAM
	Zou Zou Kuzyk	University of Manitoba
	Michaela de Melo	UQAM
	Fanny Noisette	UQAR
	Mary O'Connor	University of British Columbia
	Mila Oser (Secretary)	Hydro-Québec
<b>Absent:</b>	Daniel Brosseau	Hydro-Québec
	Jean-Philippe Gilbert	Hydro-Québec
	Graeme Morin	Cree nation Government
	Roderick Pachano	Cree Nation of Chisasibi

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## MEETING CHAIR AND SECRETARY

Marc Dunn chaired the meeting, and Mila Oser acted as the meeting secretary.

## PROPOSED AGENDA

1. Approval of the Agenda
2. Approval of the minutes from the previous meeting – December 4, 2024
3. Presentation of research paper "Riverine influence on physicochemical properties of coastal waters along a latitudinal gradient in the eastern James Bay
4. Update of the Landscape Change project of Phase II
5. Presentation of research paper "Winter to summer transition in seawater salinity, temperature, and light at Eelgrass Bed Habitats in northeastern James Bay"
6. CHCRP Eelgrass Team Final Report 2022-2024
7. Miscellaneous
8. Summary and Next Steps
9. Next Meeting

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### 1. Approval of the Agenda

The Chair reviewed the agenda, and no additional points were proposed. Thus, the agenda was approved as presented

### 2. Approval of the minutes from the previous meetings

The Chair and Luc Duquette (Mr. Duquette) discussed the approval of the minutes from the previous meeting held on December 4, 2024. They emphasized that until the minutes are officially approved, they should not be shared with non-committee members. This is to prevent the dissemination of potentially inaccurate or incomplete information. The minutes were approved.

### 3. Presentation of research paper "Riverine influence on physicochemical properties of coastal waters along a latitudinal gradient in the eastern James Bay"

Paul del Giorgio (Mr. del Giorgio) delivered a presentation titled "Riverine influence on physicochemical properties of coastal waters along a latitudinal gradient in the eastern James Bay, and a copy of the presentation and document is appended to these minutes for reference.

Mr. del Giorgio acknowledged the support and collaboration from various stakeholders, including land users, communities, Hydro Québec, Niskamoon, and fellow researchers. The presentation highlighted the importance of partnerships and the ongoing collaboration that has enabled this research. Mr. del Giorgio introduced his colleague, Michaela de Melo from UQAM, who also worked on this research and manuscript and was joining the meeting in person.

Mr. del Georgio presented an update on a manuscript that has been submitted for publication. The manuscript focuses on the river team's collaboration with the coastal team to understand the influence of rivers on the coastal water quality in the James Bay area. The study leverages datasets collected during Phase 1 of the Coastal Habitat Comprehensive Research Project (CHCRP) to determine the links between river and coastal properties. The main aim is to assess how rivers influence the coastal system beyond their immediate plumes, addressing a complex question that has not been widely studied globally.

Key findings include:

- Rivers in the James Bay area cluster based on their water quality, with southern rivers being murkier and northern rivers having more organic matter.
- The coastal sites also cluster based on water quality, and these clusters align with the river clusters, indicating that rivers influence coastal water quality along the entire James Bay.
- The river La Grande – Chisasibi stands out as it has a unique water quality profile, different from other rivers in the area.

Mr. del Georgio said that the manuscript concludes that rivers significantly influence the coastal water quality along the entire James Bay, not just at the plume but also beyond it. This finding is important as it suggests that changes in river conditions due to landscape or climate changes can impact the coastal water quality.

Mr. Duquette asked for clarification on what a cluster of coastal sites means in the context of the presentation. Mr. del Georgio explained that the sites cluster based on a set of environmental variables and that this clustering is statistical rather than geographical.

Mr. Dunn mentioned that although he had not seen the manuscript yet, he wanted to make sure it included a Cree perspective, as the Crees have long recognized the link between the coast and rivers. This paper should mention their acknowledgment of this connection.

Mrs. de Melo said that during Phase 1, they installed hydrometric stations, which are still operating. In winter, ice causes problems, and adjustments are made in spring, resulting in some data gaps. This data is downloaded weekly and uploaded to the website for everyone to view. Mrs. de Melo also shared the link to the Eeyou River discharge [website](#) explaining its format and the available data.

In relation to data collection along the river, Mr. Duquette also mentioned that following a conversation with the geomatic team the day before, some concerns were raised about the hydrometric stations along the river and their maintenance. One of the main concerns was the need for regular maintenance and calibration of the hydrometric stations to ensure accurate data collection. Mr. Dunn suggested that a dedicated team should be responsible for the upkeep of these stations to prevent any discrepancies in the data.

Additionally, there was a discussion about the accessibility of the data collected by these stations. And after discussions among the members of the steering committee members, the creation of a centralized database where all the data from the hydrometric stations could be stored and accessed easily by

researchers and stakeholders was proposed. Another suggestion was to enhance the existing hydrometric stations with advanced sensors and technology to improve the quality and range of data collected. This would help in better understanding the river's behavior and its impact on the surrounding environment.

#### 4. Update of the Landscape Change project of Phase II

*It is noted that Mr. del Georgio covered Phase 2 in the above part of his presentation*

Mr. del Georgio said that during CHCRP Phase 2, the focus was on understanding the influence of rivers on coastal water quality in the James Bay area. The study aimed to determine the links between river and coastal properties, addressing the complex question of how rivers influence the coastal system beyond their immediate plumes.

Key findings include:

- Rivers significantly impact coastal water quality along the entire James Bay, not just at the plume but also beyond it.
- Southern rivers are murkier with more sediments, while northern rivers have more organic matter.
- The river La Grande has a unique water quality profile, different from other rivers in the area.

The study leveraged datasets collected during CHCRP Phase 1 to determine the links between river and coastal properties. The findings suggest that changes in river conditions due to landscape or climate changes can impact the coastal water quality.

Mr. del Georgio mentioned that they were hoping to submit the Alliance NSERC proposal during the course of the winter and get a reply by this summer.

#### 5. Presentation of research paper "Winter to summer transition in seawater salinity, temperature, and light at eelgrass bed habitats in northeastern James Bay"

Zou Zou Kuzyk (**Mrs. Kuzyk**) gave an update of the research and findings from the document «Winter to summer transition in seawater salinity, temperature, and light at eelgrass bed habitats in northeastern James Bay that was already circulated by Mrs. Leblanc.

Mrs. Kuzyk explained that in early 2019, during the winter, they collaborated with community members and were able to deploy instruments in March 2019, with a second deployment on April 4. They placed the instruments at the bottom near eelgrass beds because we were concerned about ice, which could cause jams in the spring, so they positioned the instruments deeper to avoid losing them.

They remained deployed from March until the following August, and the data presented in this document are from that period.

Mrs. Kuzyk said that one of the main things they learned in CHCRP phase 1 is that the conditions in the spring seem to be particularly important.

Mrs. Kuzyk mentioned the take home messages:

Eelgrass requires a significant amount of light, and their studies indicate that neither of the sites they examined provide the maximum light needed for eelgrass. At this stage, Mrs. Kuzyk said that they are unable to develop a universal model that provides exact outcomes, but they have gained substantial knowledge and established a solid foundation.

During the meeting, there was a discussion about the feedback on the CHCRP Phase 1 synthesis document, which had been circulated by Mrs. Leblanc. Mrs. Kuzyk provided comments on this document, highlighting the need for a more comprehensive synthesis of the research findings.

In response to Mrs. Kuzyk's comments, it was agreed that the synthesis document would be revised to incorporate her feedback. The steering committee team decided to work on improving the document to ensure it accurately reflects the research findings and addresses the points raised by Mrs. Kuzyk.

## **6. CHCRP Eelgrass Team Final Report 2022-2024**

Fanny Noisette, (Mrs. Noisette) delivered a presentation titled "Presenting CHCRP Phase II (2023-2024)," and a copy of the presentation and document is appended to these minutes for reference.

Mrs. Noisette gave a reminder of the timeline: phase 1 (2019-2021), the interim phase (2022-2024), and phase 2 (2024-2029) and said this is a joint effort between stakeholders.

Mrs. Noisette mentioned that eelgrass persists but has not fully recovered.

Mrs. Noisette mentioned that the most important aspect of CHCRP phase 1 was reported in the final eelgrass report in phase 1.

One of the objectives of the interim phase was to follow eelgrass growth in five locations to understand the connection with early summer conditions.

Mrs. Noisette mentioned that between these interim phases, several events occurred, such as the forest fires in the summer of 2023, which put the study on standby. During this time, the team piloted new methods focusing on rhizome growth and sugar content. Fieldwork took place in Eeyou Istchee in summer 2024.

Mrs. Noisette said that their findings showed that although more sugar coincides with more growth, the sugar itself was not directly correlated to growth. As rhizomes grow, sugar is present simultaneously, and this helped them understand how eelgrass can grow during winter.

In the summer interim phase, from the end of June to July 2024, Mrs. Noisette said they conducted samplings and held community meetings to discuss the project and priorities and continued monitoring through additional samplings.

Assessment of eelgrass condition in key locations:

Mrs. Noisette said that eelgrass shoot lengths, density, and biomass were comparable to regional averages in recent years but generally remain below pre-decline levels.

The researchers followed eelgrass growth in five locations in the Chisasibi and Wemindji areas to understand the link between early summer growth and environmental conditions. Shoots were pricked to monitor and measure growth over ten days. At the CH3 site, growth was faster. They also observed a significant number of algae mats in several locations this year, which could affect eelgrass growth.

Mrs. Noisette mentioned that Niskamoon requires this report to be approved to release the final payment.

## **7. Miscellaneous**

Following a request from Chief Daisy House to share steering committee documents with the CTA and CERRI, extend an invitation to Dante Torio (CERRI) as an observer at the next steering committee meeting.

## **8. Summary and Next Steps**

Approval of Minutes: Ensure that the minutes from the previous meeting (December 4, 2024) are approved and not shared with non-committee members until they are officially approved.

Video Script and Narration Review: Approve the video script and narration review comments in January.

Data Finalization: Finalize the data from the Long Island project.

Synthesis Document Revision: Revise the CHCRP Phase 1 synthesis document to incorporate feedback from Mrs. Kuzyk.

Review the manuscript submitted by Michaela de Melo and Paul del Georgio and give feedback to Mila Oser by February 20<sup>th</sup>.

## **9. Next Meeting**

Following the exchange on the availability of each, it was agreed that the next meeting will be held on Friday, February 28, 2025, from 9:00 AM to 12:00 PM, via Teams.

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**ADJOURNMENT OF THE MEETING**

Considering that all items on the agenda were addressed, the meeting is adjourned at 12:30 PM.

The meeting secretary,

*Mila Oser*

Mila Oser

The meeting Chair,

A handwritten signature in black ink, appearing to read 'Marc Dunn', with a horizontal line extending to the right.

Marc Dunn

1 **Riverine influence on physicochemical properties of coastal waters along a**  
2 **latitudinal gradient in the eastern James Bay**

3 Michaela L. de Melo<sup>1\*</sup>, Caroline Fink-Mercier<sup>2</sup>, Virginie Galindo<sup>2</sup>, Michel Gosselin<sup>2</sup>, Urs  
4 Neumeier<sup>2</sup>, Huixiang Xie<sup>2</sup>, and Paul A. del Giorgio<sup>1</sup>

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10 ***Abstract***

11 Rivers integrate climate signals, landscape gradients and environmental disturbances at the  
12 watershed scale, motivating the effort to better understand the influence of riverine exports on  
13 downstream ecosystems. Here we aim to establish connections between the physicochemical  
14 properties of rivers draining into the eastern James Bay (JB), and of the coastal waters along its  
15 entire eastern shore. We clustered 17 river outlets and over 140 coastal sites along the latitudinal  
16 gradient (~300 km shoreline) of JB sampled during two consecutive summers according to trends  
17 in nutrients, suspended particulate matter (SPM), colored dissolved organic matter (CDOM)  
18 absorbance, freshwater discharge and salinity. The transition zones, where significant latitudinal  
19 changes in water physicochemical variables occur, were generally spatially consistent between  
20 rivers and coastal waters. Average material concentrations were overall higher in rivers than in  
21 adjacent coastal waters. We conclude that rivers broadly shape the coastal physicochemical

22 conditions along eastern JB, and that the riverine influence varies among areas as a function of  
23 streamflow and of the variable considered.

24 Keywords: CDOM, turbidity, nitrogen, phosphorus, cluster analysis, transition zones

25 Running head: Riverine influence on eastern James Bay coast

26

### 27 *Scientific Significance Statement*

28 The subarctic James Bay is the most river-affected water body of the Hudson Bay system and  
29 represents a major compartment to study riverine influence on coastal biogeochemistry. Although  
30 freshwater exports to the James Bay have previously been quantified, we still lack an integrated  
31 perspective on how river exports affect coastal biogeochemistry. This study identified transition  
32 zones of salinity, nutrients, turbidity, and colored dissolved organic matter (CDOM) absorbance  
33 along the coast of the eastern James Bay and estimated river influence on coastal waters in distinct  
34 areas during summer as a function of streamflow and variable considered.

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## 42           **Introduction**

43           Rivers are major conduits for the transport and processing of dissolved and particulate  
44 materials (e.g., nutrients and suspended particles) from land to the ocean. By integrating climate  
45 signals (e.g., precipitation), watershed properties (e.g., land cover) and environmental disturbances  
46 (e.g., wildfires), rivers act as sentinels of landscape and climatic changes and influence  
47 downstream ecosystems through export of terrestrial organic matter, sediments, contaminants, and  
48 macro- and micronutrients (Dupas et al., 2017; Li Yung Lung et al., 2018). Such exports are critical  
49 to biogeochemical cycling and biological productivity in coastal ecosystems (Grimes 2001;  
50 Terhaar et al. 2021). Research to date has focused on characterizing seasonal and long-term  
51 dynamics of individual river watersheds, as well as identifying major underlying drivers of  
52 material concentration and export (Tank et al. 2012; McClelland et al. 2014; Moquet et al. 2016;  
53 Li Yung Lung et al. 2018), but less attention has been paid on how these riverine exports shape  
54 coastal material budgets.

55           It has been well established that rivers exert strong local influence along their respective  
56 coastal plumes, and there has been extensive research on the processes that mediate the export of  
57 materials to open oceans, such as flocculation, deposition, and other biogeochemical processes  
58 (Emmerton et al. 2008; Bauer et al. 2013; Asmala et al. 2017). However, few studies have assessed  
59 the extent to which broad hydrologic and environmental gradients involving multiple watersheds  
60 at regional scales may generate patterns in physical and biogeochemical properties in marine  
61 coastal areas over large spatial scales (Zhang and Blomquist 2018; Frigstad et al. 2020).

62           This is certainly the case for understudied regions of the Canadian Arctic (Li Yung Lung  
63 et al. 2018), where freshwater exports are an important fraction of the nearshore marine volume  
64 (McClelland et al. 2012). Located in the southern end of Hudson Bay, the James Bay (JB) is the

65 most river-influenced water body of the Hudson Bay system (Déry et al. 2005), receiving a river  
66 runoff from its eastern shore of about 227 km<sup>3</sup> per year (de Melo et al. 2022). Regional hydrology  
67 has been significantly altered by river diversions (about 65 km<sup>3</sup> per year) and damming for  
68 hydropower production, and there are large climatic and environmental gradients within the  
69 territory that result in a wide range of watershed landscape properties (de Melo et al. 2022). It has  
70 been shown that hydrologic export from these watersheds directly shape river plumes and  
71 surrounding coastal habitats in the JB, influencing local salinity, turbidity, nutrients, and dissolved  
72 organic matter (Peck et al. 2022; Évrard et al. 2023; Kuzyk et al. 2023; Leblanc et al. 2023;  
73 Meilleur et al. 2023). Beyond these localized riverine effects, there is also evidence of larger scale  
74 latitudinal physicochemical gradients (Évrard et al. 2023; Lee et al. 2023; Guzzi et al. 2024) along  
75 the entire eastern JB coast, yet the underlying drivers of these patterns are unclear, and so is the  
76 potential contribution of regional rivers to these patterns.

77         The objective of this study is to assess potential links between broad gradients in river  
78 properties and large-scale physicochemical patterns along the eastern JB, mediated by river  
79 transport of water and materials from land to the coast. This exploration is fundamental because  
80 the terrestrial landscape and the hydrology of the region are drastically changing (Royer and  
81 Herrmann 2013; Déry et al. 2016; de Melo et al. 2022), but it is uncertain how these shifts will  
82 impact the biogeochemical functioning of the JB beyond very localized river plume effects. Yet,  
83 there are potential far-reaching implications of these impacts on coastal productivity and  
84 biodiversity. The aim is to determine large-scale physicochemical gradients and identifying  
85 discontinuities in relevant variables, such as turbidity, nutrients, and CDOM absorbance, in both  
86 rivers and coastal habitats at comparable spatial and temporal scales along the entire eastern JB,  
87 and to subsequently explore potential connections. Here, we identified transition zones where

88 significant changes in both riverine and coastal water physicochemical properties occur following  
89 the regional ocean circulation pattern, and we have assessed the degree to which these transitions  
90 in rivers and coastal sites overlap along the Bay.

91

## 92 **Methods**

### 93 **Study area**

94 This study was conducted along the eastern coast of James Bay (JB), the southern extension  
95 of Hudson Bay in northern Quebec, Canada, within the traditional Cree territory (Eeyou Istchee),  
96 as part of the Coastal Habitat Comprehensive Research Project (CHCRP, Kuzyk et al. 2023), an  
97 interdisciplinary, Cree-driven community-academic partnership.

98 JB is characterized by a cyclonic circulation during the open water season (Prinsenber  
99 1986) and a moist continental subarctic climate (Koeppen Climate Classification System) with  
100 contrasting seasons: ice-covered cold winters and warm to cool summers, with a short growth  
101 season from June to November (Davis et al. 2024). Ice formation in the Hudson Bay system  
102 progresses from northwest to southeast, with JB ice-covered by early December. Ice breakup starts  
103 in late May or early June near river mouths and marine inflows, and JB is usually ice-free by early  
104 July (Taha et al. 2019).

105 The eastern JB drainage basin lies predominantly within the Canadian Shield  
106 physiographic region, shaped largely by glaciation, and consists mainly of a coastal plain with  
107 many lakes, ponds, peat bogs, and swamps (The Atlas of Canada, Canada 1978). The Great Clay  
108 Belt, the largest clay pocket within the Canadian Shield (Dresser 1913), extends across watersheds  
109 in the southern region of the Bay, with several rivers flowing into Rupert and Hannah bays. Several

110 river watersheds draining an area over 350 000 km<sup>2</sup> discharge a total water volume of 227 km<sup>3</sup>  
111 yearly into the eastern JB (de Melo et al. 2022). The La Grande (LG) River contributes to over  
112 50% of total freshwater inputs to the JB during all seasons due to several watershed diversions for  
113 hydropower production in the past decades, resulting in over doubling its natural mean annual  
114 discharge and inverting its seasonal flow pattern with now the discharge peaking in winter (Déry  
115 et al. 2016).

116

### 117 **River and coastal habitat sampling**

118 A total of 17 rivers were sampled close to their mouth from the Harricana River in the  
119 south (48°34'13" N, 78°07'17" W) to the Salmon River in the north (54°33'35" N, 79°25'05" W)  
120 covering more than 400 km of coastline (Figs. 1A and S1). They were accessed by boat or  
121 helicopter in summer 2018 (July 17-27) and 2019 (July 31 - August 13). Coastal marine sites were  
122 sampled from the Jolicoeur River (~52° N) to near the mouth of the Piagochioui River (~54° N),  
123 along approximately 250 km of coastline, by freighter canoes guided by Cree guides (Fig. 2A).  
124 The coastal team sampled over 140 sites in summer 2018 (August 01-18) and 2019 (July 05 –  
125 August 14), providing a robust latitudinal coverage of the coast (Fig. S1; comparison between  
126 years in Table S1).

127

### 128 **Environmental variables**

129 Water samples were collected at depth of 0.5 m using a peristaltic pump and stored in 20  
130 L acid washed polycarbonate containers (river) or using a 5 L Niskin bottles (coast). Samples were  
131 then properly filtered (for dissolved fractions) and stored in pre-cleaned polypropylene specimen

132 containers for nutrients and in glass bottles for CDOM absorbance measurements within hours  
133 before being transported for analytical laboratories. The total nitrogen concentration (TN,  
134 unfiltered water sample) of river samples was analyzed in duplicate using the alkaline persulfate  
135 digestion in an Alpkem Flow-Solution IV autoanalyzer (O I Analytical, College Station, TX,  
136 USA). For dissolved nitrogen concentration (DN), coastal water was filtered onto a pre-combusted  
137 Whatman GF/F filter (450°C for 5 h, 25 mm, nominal porosity of 0.7  $\mu\text{m}$ ) and analyzed in a  
138 Shimadzu TOC- $V_{\text{CPN}}$  analyzer (Kyoto, Japan) with a chemiluminescent nitrogen detector (TNM-  
139 1 module). Here, TN from rivers is compared to DN from coastal sites because no reliable TN data  
140 was available for coastal sites. The dissolved fraction has been shown to be the dominant fraction  
141 in coastal surface waters (Guo et al. 2004; Duan et al. 2016) and possibly accounted for 99% (sd  
142 = 12%) of TN in the present study (Fig. S2).

143 The total phosphorus (TP) concentration in river samples was determined in duplicate after  
144 persulfate digestion as orthophosphate with the molybdenum blue spectrophotometric technique  
145 (890 nm, Ultrospec 2100 pro, Biochrom Ltd., Cambridge, UK), while dissolved inorganic  
146 phosphorus samples (i.e., phosphate,  $\text{PO}_4^{3-}$ ) from the coast were filtered onto pre-combusted  
147 Whatman GF/F filters and analyzed using a Bran-Luebbe autoanalyzer 3 (Bran+Luebbe GmbH,  
148 Norderstedt, Germany - adapted method from Hansen and Koroleff 1999).

149 For determination of CDOM absorbance, water was filtered through 0.45 or 0.2  $\mu\text{m}$   
150 polyethersulfone membranes for river and coastal waters, respectively and stored in pre-cleaned  
151 glass flasks in 4 °C in the dark until analysis within one month of sample collection (for coast, see  
152 Évrard et al. 2023) and no more than 2 months (for rivers). Previous studies have shown negligible  
153 differences when using filters of different pore sizes for CDOM absorbance measurements in  
154 waters with significant terrestrial inputs (0.7 and 0.2  $\mu\text{m}$ , Massicotte et al. 2017). Samples were

155 scanned over 200-800 nm using an Ultrospec3100 pro spectrophotometer for riverine samples  
156 (Biochrom) and a Lambda-35 dual beam UV–visible spectrophotometer (PerkinElmer, Waltham,  
157 MA, USA) for coastal samples, both fitted with 1 or 5 cm quartz cells and referenced to nanopure  
158 water. A baseline correction was applied by subtracting the absorbance at 690 nm for river samples  
159 or the average absorbance value between 683 and 687 nm for coastal samples (680-690 nm, Babin  
160 et al. 2003). The Napierian absorption coefficient at 440 nm ( $a_{CDOM(440)}$  ( $m^{-1}$ )), calculated as  
161 2.303 times the absorbance at 440 nm divided by the cell’s pathlength in meters, was chosen as an  
162 indicator of CDOM abundance in the eastern JB (Mabit et al. 2022; Évrard et al. 2023).

163         Suspended particulate matter (SPM) was determined by filtering water samples onto pre-  
164 combusted and pre-weighed 0.7  $\mu m$  Whatman GF/F filters in duplicate for river sites or triplicate  
165 for coastal sites. Filters were dried at 80°C for 24 h and weighed on an analytical balance (adapted  
166 method from Neukermans et al. 2012). We averaged the dry weights of the material collected in  
167 the replicates.

168         Chlorophyll *a* (chl *a*) in river samples were filtered in duplicate on Whatman GF/F filters  
169 and extracted with 90% hot ethanol in the dark. Chl *a* concentrations were determined  
170 spectrophotometrically in ethanol extracts (Wintermans and De Mots 1965). A turbidity correction  
171 (665 nm-750 nm) and phaeophytin correction after acidification (0.01 N HCl final concentration)  
172 were performed (Lorenzen 1967; Nusch 1980). Concentrations from coastal sites were measured  
173 using a Turner Designs AU-10 fluorometer (San Jose, CA, USA) after 24 h extraction in 90%  
174 acetone at 4° C in the dark (acidification method, Parsons et al. 1984).

175         Summer freshwater discharge (from July to September) was determined using high-  
176 frequency data from autonomous hydrometric stations installed in river mouths as described  
177 previously (de Melo et al. 2022). Salinity was measured at the coastal sites as conductivity using

178 a SBE 19plus V2 CTD (Sea-Bird Scientific, Bellevue, WA, USA) and with a Cond 330i  
179 conductivity meter (WTW, Weilheim, Germany) calibrated against a KCl solution (1413  $\mu\text{S cm}^{-1}$ ), both validated with discrete samples analysed in the laboratory, with an 8410A Portasal  
180 salinometer (Guildline Instruments, Smiths Falls, CA, USA). Both approaches for determining the  
181 salinity agreed well (Évrard et al. 2023).

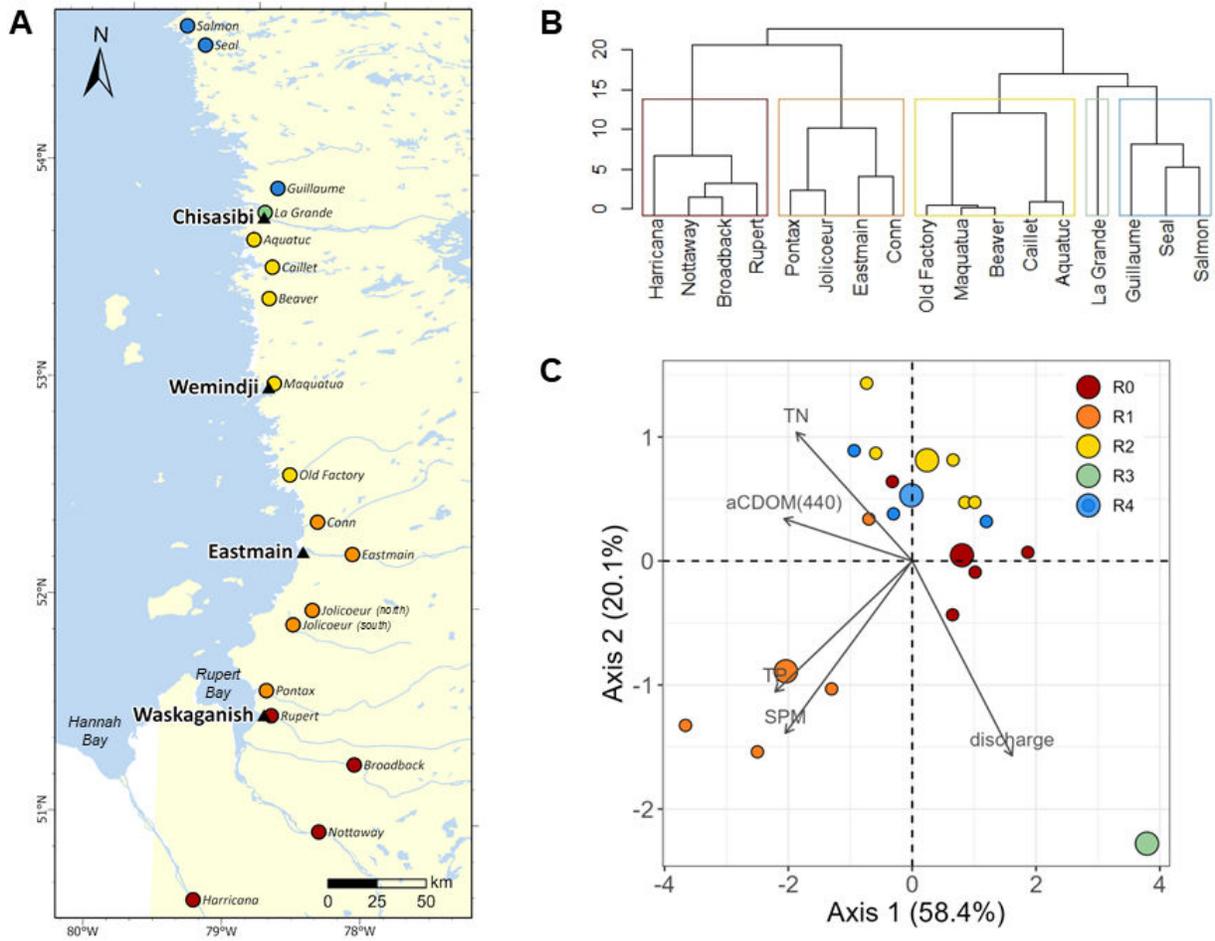
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## 184 **Statistical analyses**

185 The variables retained to perform the clustering analyses were selected by their potential  
186 influence on the coastal environment: 1) light penetration ( $a_{\text{CDOM}(440)}$  and SPM); 2) nutrient  
187 concentrations (nitrogen – DN and TN for coastal and river sites, respectively; phosphorus –  $\text{PO}_4^{3-}$   
188 and TP for coastal and river sites, respectively); and 3) water salinity (coastal sites) or summer  
189 freshwater total discharge (rivers). Hierarchical clustering constrained by latitude order (of  
190 Euclidean distances of normalized data matrices, Camêlo Aguiar et al. 2020) was used to group  
191 rivers and coastal sites, separately, based on the similarity of their physicochemical variables.  
192 Choosing the optimal number of clusters is subjective due to the unsupervised nature of clustering  
193 (Liu et al. 2022). Here, we checked the total within sum of squares and compared the dispersion  
194 of our hierarchical classifications to that obtained from a broken stick model (bstick function) to  
195 guide the choice of cluster groups (Bennett 1996, Fig. S3). For hierarchical clustering, we first  
196 ordered sampling sites following a latitudinal gradient from south to north, we then normalized  
197 the data using the function `decostand` (method= “standardize”) in R (version 4.0.2) and computed  
198 Euclidean distance matrices using the function `vegdist()`. Then, we performed a constrained  
199 hierarchical clustering of the distance matrices using the function `chclust()` (package “rioja”) with  
200 clusters constrained by latitudinal order. Principal component analysis (PCA) was used to

201 summarize differences among clusters based on physicochemical variables and to assess the  
202 percentage of variation explained.

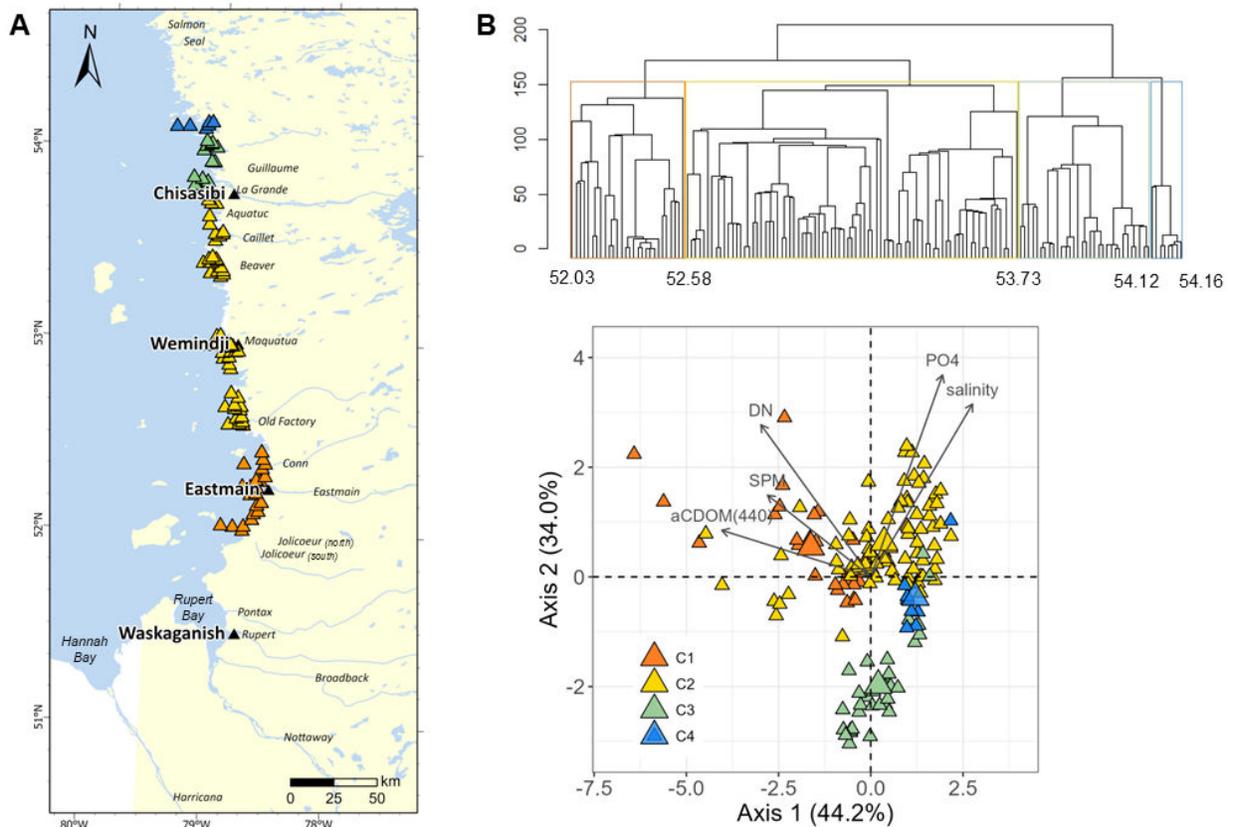
203         We tested for differences in physicochemical variables between clusters within rivers and  
204 coastal sites using One-way ANOVA or Kruskal-Wallis tests followed by post-hoc Dunn's or  
205 Tukey tests (after assessing for data normality and homoscedasticity). Due to the small number of  
206 river samples per cluster ( $n = 1$  to  $5$ ), we used values from the two summers separately (while  
207 maintaining the same cluster classification) to achieve more robust statistics ( $n = 2$  to  $10$ ). We  
208 calculated relative differences between the average observed concentration of each variable  
209 (nutrients (i.e., TN vs. DN and TP vs.  $\text{PO}_4$ ),  $a_{\text{CDOM}}(440)$ , and SPM) measured in river clusters and  
210 the corresponding coastal cluster as a proxy for the degree of river influence. Differences in mean  
211 concentrations were divided by river mean concentrations, and then multiplied by 100 to obtain  
212 percentages of riverine decoupling. Negative relative differences indicate higher values at coastal  
213 sites than at river sites. All analyses were conducted in R and maps were created using ArcGIS 10.



214

215 **Figure 1.** (A) Map of the sampled rivers with dot colors representing river clusters R0 to R4; (B)  
 216 Dendrogram of the hierarchical clustering based on physicochemical variables (TN, TP,  
 217  $a_{CDOM}(440)$ , SPM, and summer freshwater discharge) and constrained by latitudinal order of river  
 218 sites sampled in summers 2018 and 2019; (C) Principal component analysis (PCA) of clusters  
 219 based on the same physicochemical variables as in panel (B), with larger dots representing the  
 220 centroid of each cluster. Colours used in dots in panels (A) and (C) and rectangles in (B) indicate  
 221 different clusters.

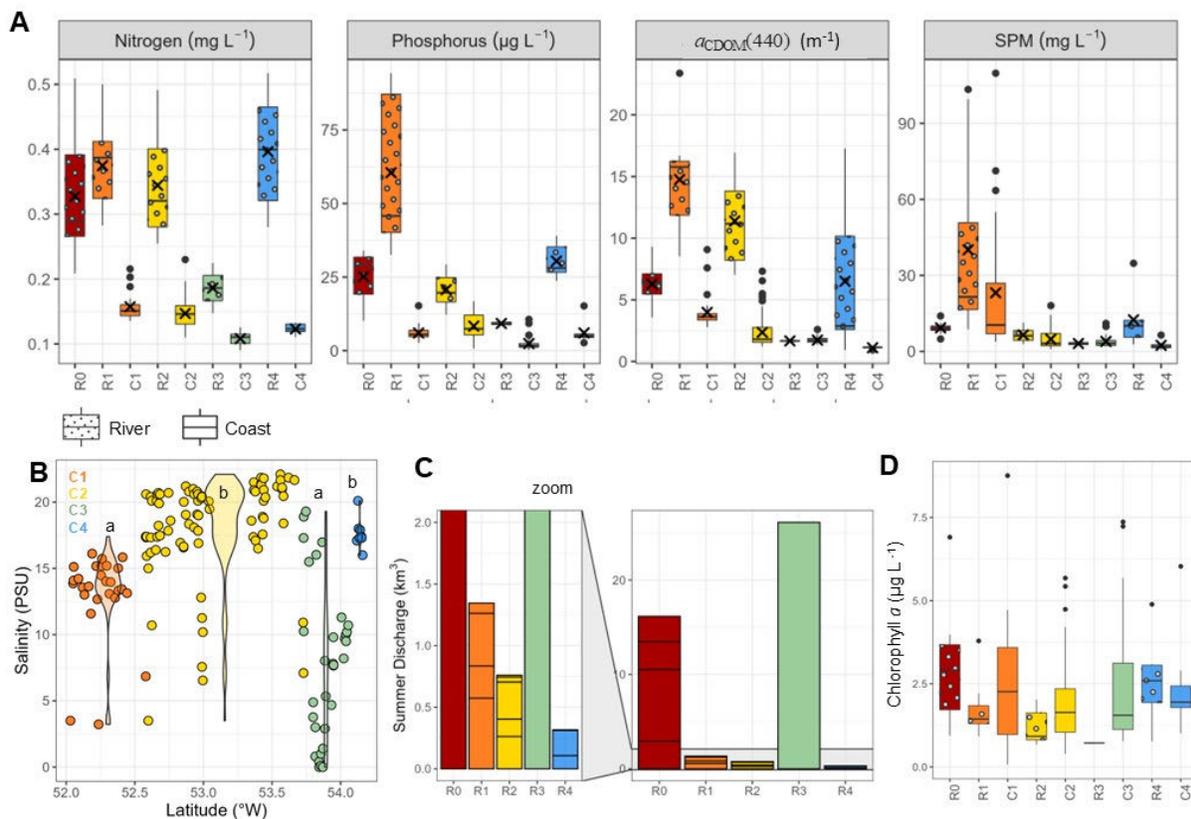
222



223  
 224 **Figure 2.** (A) Map of the sampled coastal sites with triangle colors representing clusters C1 to C4;  
 225 (B) Dendrogram of the hierarchical clustering based on physicochemical variables (DN,  $\text{PO}_4^{3-}$ ,  
 226  $a_{\text{CDOM}}(440)$ , SPM, and salinity) and constrained by latitudinal order of coastal sites sampled in  
 227 summers 2018 and 2019; (C) Principal component analysis (PCA) of clusters based on the same  
 228 physicochemical variables as in panel (B), with larger triangles representing the centroid of each  
 229 cluster. Colours used in triangles in panels (A) and (C) and rectangles in (B) indicate different  
 230 clusters.

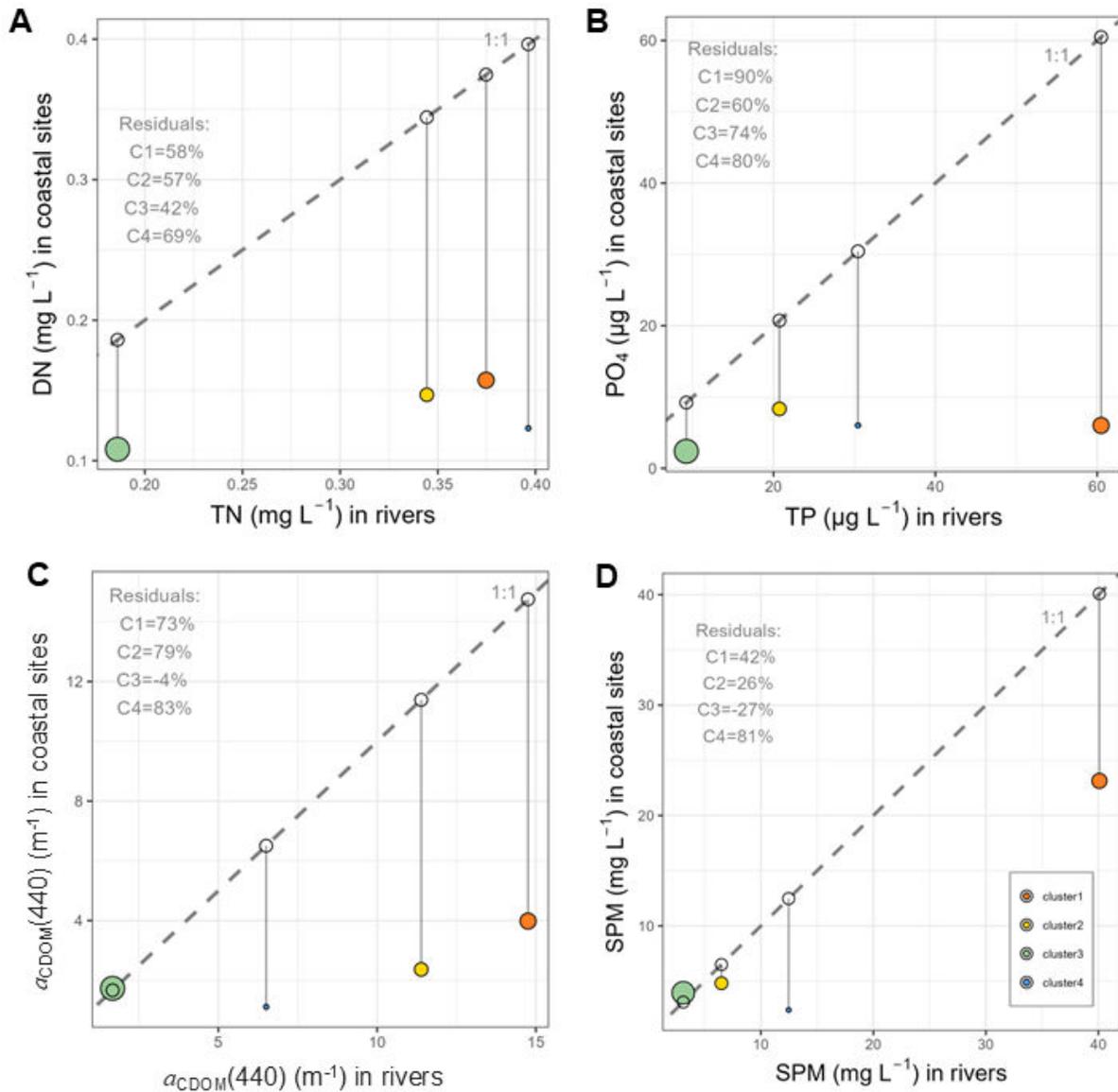
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232



233  
 234 **Figure 3.** (A) Boxplots comparing nitrogen (TN or DN), phosphorus (TP or  $\text{PO}_4$ ),  $a_{\text{CDOM}(440)}$  and  
 235 SPM values at river (R) and coastal (C) sites. Horizontal lines represent the median and x the mean,  
 236 first and third quartiles, and whiskers extend to the furthest data point within 1.5 times the  
 237 interquartile range; statistics are presented in Table S3 and S4. Colours and textures represent  
 238 cluster and system types (i.e., river vs coastal sites), respectively. (B) Salinity at coastal sites along  
 239 the latitudinal gradient, where dots represent individual values and violin plots represent variation  
 240 within each coastal cluster (represented by colours), letters *a* and *b* represent significant differences  
 241 in salinity between clusters (Kruskal-Wallis - Table S4); (C) Histogram representing total summer  
 242 freshwater discharge per cluster. Segments within bars represent the contribution of each river  
 243 individually. (D) Chlorophyll *a* concentration in river and coastal clusters.

244



245

246 **Figure 4.** Relative differences between mean (A) nitrogen concentrations, (B) phosphorus  
 247 concentrations, (C)  $a_{\text{CDOM}}(440)$  and (D) SPM in each river cluster, and its corresponding coastal  
 248 cluster (i.e., R1-C1). Small differences (close to the line 1:1) represent a strong river influence on  
 249 the coastal waters. Values above the 1:1 line (also presented as negative % in the text box)  
 250 represent higher mean values in coastal waters than in rivers, and values below the 1:1 line  
 251 represent lower mean values in coastal waters than in rivers (positive % in the text box). Dot colors

252 represent the clusters, while sizes represent the sum of river discharge for each cluster (log 10  
253 transformed).

254

## 255 **Results**

### 256 **Latitudinal clustering of rivers and coastal sites**

257 Variation in river physicochemical variables (Table S2) were captured by hierarchical  
258 clustering analysis, which suggests the existence of five distinct river clusters along the latitudinal  
259 gradient from south to north (Fig. 1A, B). More specifically, river cluster 0 (R0) includes the rivers  
260 from Harricana to Rupert, R1 includes the rivers from Pontax to Conn, R2 includes the rivers from  
261 Old Factory to Aquatuc, R3 is represented exclusively by La Grande (LG) River, and R4 includes  
262 Guillaume, Seal, and Salmon rivers. The variability in biogeochemical conditions among rivers  
263 was captured along two principal axes in the PCA, which collectively explained 78.5% of the  
264 variance in their physicochemical properties (Fig. 1C). Axis 1 (i.e., 58.4% of the total variance)  
265 primarily reflected differences in  $a_{CDOM}(440)$  and TP, whereas axis 2 (20.1%) was predominantly  
266 related to TN and summer freshwater discharge. SPM was similarly related to both axes.

267 Coastal sites (Fig. 2A) clustered into four distinct groups and followed a very similar  
268 latitudinal pattern to the one observed for rivers (Figs. 1B and 2B). Distinct transition zones were  
269 observed at latitudes 52.03°N (C1), 52.58°N (C2), 53.73°N (C3), and ca. 54.1°N (C4). The two  
270 main PCA axes explained 78.2% of the variance in physicochemical variables collected at the  
271 coastal sites. Axis 1 (i.e., 44.2% of the total variance) was mainly related to  $a_{CDOM}(440)$  and  
272 secondarily to SPM, whereas axis 2 (34%) was mainly related to PO<sub>4</sub>, salinity, and DN.

273 Figure 3 compares the average physicochemical properties of the river (R) and coastal (C)  
274 clusters, as well as average river discharge and coastal salinity. Cluster R0 represents southern  
275 rivers characterized by high summer discharge and intermediate values of TN, TP,  $a_{CDOM}(440)$ ,  
276 and SPM compared to the other rivers studied (Fig. 3A). The transition zone between river clusters  
277 R0 and R1 is characterized by a significant increase in TP and SPM concentrations and  $a_{CDOM}(440)$   
278 (p-values < 0.05, Fig. 3A, Table S3) and a decrease of total summer freshwater discharge (Fig.  
279 3C). River clusters R1 and R2 had similarly high  $a_{CDOM}(440)$ , but R1 had the highest SPM of all  
280 eastern JB rivers, likely reflecting drainage over the Canadian Great Clay Belt, and the highest TP  
281 concentrations, likely associated to these high particulate loads, whereas R2 had significantly  
282 lower average TP and SPM (p-value < 0.05) concentrations.

283 LG River itself formed a cluster (R3), mainly because it exports more than half of the total  
284 summer freshwater discharge to the eastern JB (~26 km<sup>3</sup>), and seasonally transports large amounts  
285 of freshwater and dissolved and particulate materials to the Bay, despite having lowest  
286 concentrations or  $a_{CDOM}(440)$  values per volume unit (Fig. 3A). All the material concentrations  
287 were consistently lower in the LG River, a pattern observed previously (de Melo et al. 2022),  
288 although we were unable to adequately test the significance of this pattern due to the small number  
289 of samples (n = 2). Finally, rivers in the northernmost cluster (R4) had consistently higher TN and  
290 TP concentrations and  $a_{CDOM}(440)$  relative to the LG River and intermediate values relative to  
291 southern river clusters (Fig. 3A, Table S3). TN concentrations were not consistently different  
292 among river clusters (ANOVA test, p-value > 0.05), except between R3 and R4, however, such  
293 statistical difference could be an artifact of the small number of R3 samples (Table S3).

294 The coastal site clusters were also characterized by different combinations of  
295 physicochemical variables. Coastal C1 had overall higher SPM,  $a_{CDOM}(440)$  and lower PO<sub>4</sub> (note

296 that for rivers we report TP), and salinity than C2 (p-values < 0.05, Fig. 3A, Table S4), but similar  
297 DN concentrations. The transition between C2 and C3 is related to a significant decrease in  
298 salinity, most likely associated to the greater distance to the high-discharge rivers Nottaway,  
299 Broadback, and Rupert, and a decrease in DN and PO<sub>4</sub>, but no systematic change in *a*<sub>CDOM(440)</sub>  
300 and SPM (p-values > 0.05). We observed a significant decrease in *a*<sub>CDOM(440)</sub> and an increase in  
301 PO<sub>4</sub> and salinity, but no significant changes in DN and SPM (p-values > 0.05) at sites located north  
302 of LG River plume (C4). It is worth mentioning that neither the coastal waters of the southernmost  
303 sector (C0-river) nor the northern border with Hudson Bay (C4-river) were sampled in summer,  
304 limiting the possibility to effectively determine river-coast coupling in these areas.

305         Figure 3D shows the average chl *a* concentration for each of the river and coastal clusters.  
306 There was a marked northward decrease in average river chl *a* concentration from R0 (around 2.5  
307 µg L<sup>-1</sup>) to R3, which had the lowest concentrations of all sites (< 0.7 µg L<sup>-1</sup>), with a subsequent  
308 increase in R4. This trend was slightly mirrored in the coastal sites, which showed more variable  
309 chl *a* concentration within clusters. Overall, chl *a* was generally higher in coastal sites than in  
310 rivers, except in the northernmost region (C4).

311

### 312 **River influence along the coast**

313         With few exceptions, the average material concentrations were higher in the river clusters  
314 than in their coastal counterparts. To further explore these river / coastal links, we estimated the  
315 relative difference between the average material concentration in the river versus the  
316 corresponding coastal cluster for the different biogeochemical variables (Fig. 4). Overall, coastal  
317 cluster C3 appears more closely aligned to the corresponding river cluster R3 for all four materials

318 (low % difference), consistent with the LG River exerting a strong local influence due to its very  
319 large discharge as seen in the low coastal salinities. At the other extreme, coastal zone C4 was  
320 overall the most decoupled (high % difference) from river conditions, but this sector had also the  
321 lowest total river discharge. In addition, only one of the three rivers in cluster R3 can potentially  
322 influence the coastal sites of zone C4 (Fig. 1A and Fig. 2A).

323         There was an overall positive relationship between mean riverine and coastal SPM (Fig.  
324 4D), suggesting a riverine influence on coastal turbidity throughout the eastern JB. The high  
325 positive relative differences for coastal C1 and C4 also suggest fast SPM deposition at the mouth  
326 of rivers with highest sediment loads. Conversely, the negative relative difference for coastal C3  
327 suggests a dilution of marine water turbidity in C3 by very low SPM riverine water from R3 (due  
328 to long water residence time within upstream reservoirs).

329         Likewise, there was a weak but positive relationship between river and coastal  $a_{\text{CDOM}(440)}$   
330 (Fig. 4C), suggesting a potential influence of riverine inputs on coastal waters, particularly of  
331 extremely high  $a_{\text{CDOM}(440)}$  loads in C1. High positive relative differences for C1, C2 and C4  
332 indicate significant dilution, processing, and/or flocculation of CDOM at the coastal margin.

333

## 334 **Discussion**

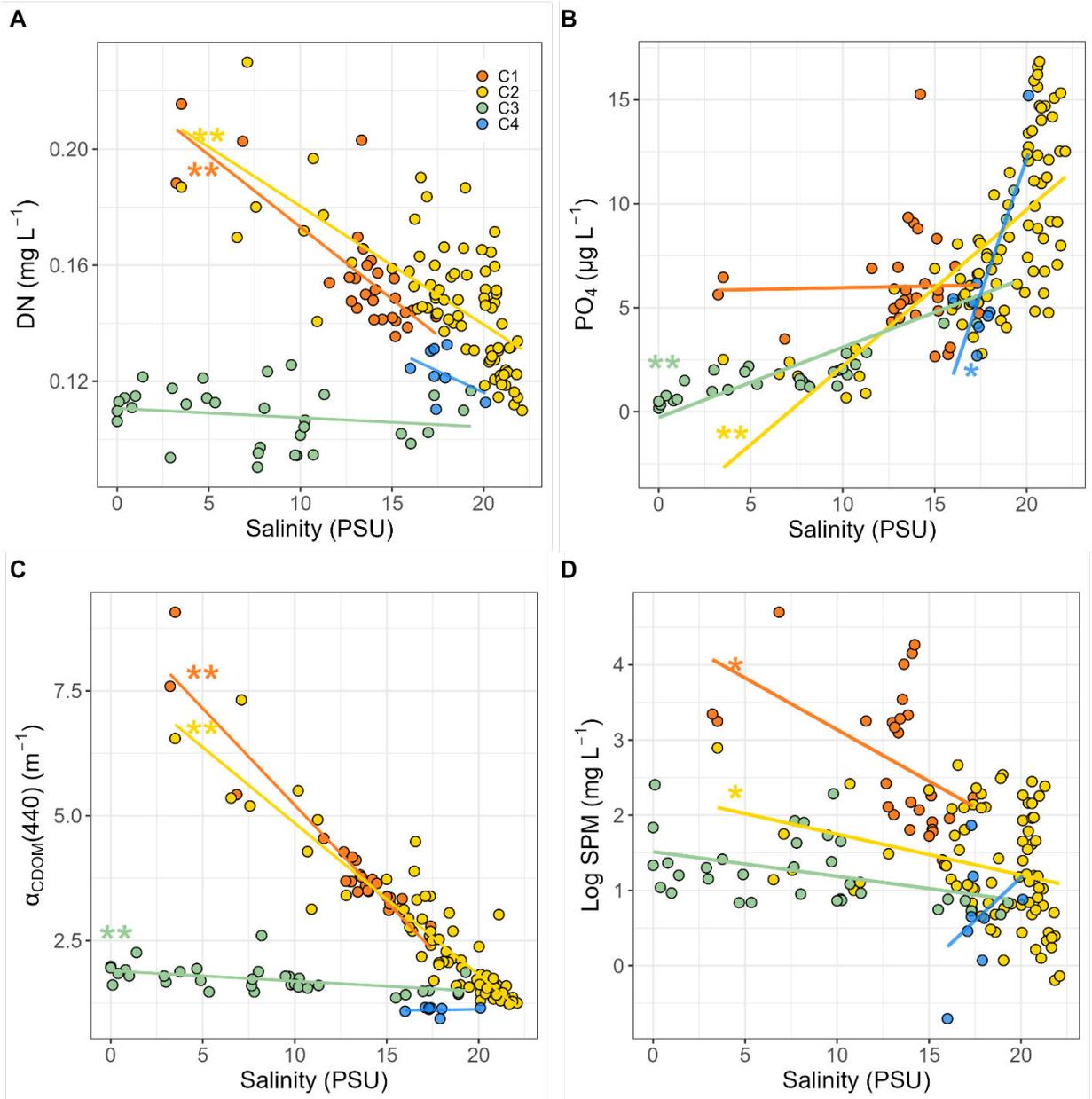
335         Our study identified significant latitudinal variability in the physicochemical properties of  
336 both rivers and coastal sites along the eastern JB, captured by distinct clusters (5 for rivers and 4  
337 along the coast). This variability resulted in transition zones (changes from one cluster to another)  
338 that were spatially coherent between river and coastal waters. For instance, the transition between  
339 the Conn and Old Factory rivers (R1-R2, C1-C2) translates into a decrease in SPM and

340  $a_{CDOM}(440)$ , and an increase in coastal salinity due to decrease in the total freshwater discharge.  
341 The highest coherence between the LG River (cluster 4) and nearby coastal waters is due to the  
342 very large freshwater export to the Bay, which strongly influences coastal conditions through the  
343 large plume extension (Peck et al. 2022; Kuzyk et al. 2023). This results in consistently low  
344 material concentrations and low salinity, in contrast to the overall Bay conditions (Évrard et al.  
345 2023).

346         Conversely, when rivers carry higher coloured and turbid waters, as in C1 and C3, this  
347 translates into coastal waters with somewhat higher colour and turbidity as well. This coherence  
348 in the location of transition zones and decreasing or increasing (or constant) pattern of material  
349 concentrations between river-coast clusters suggests a widespread influence of rivers on coastal  
350 biogeochemistry, consistent with previous studies describing the highly river-influenced nature of  
351 JB (Déry et al. 2005; Évrard et al. 2023). Other studies have emphasized that land cover and  
352 climate change strongly influence river exports to the coast, and consequently limit light  
353 penetration in the Skagerrak Sea along the Norwegian coastline. These patterns were associated  
354 with riverine CDOM and with the long-term (centennial) increase in vegetation cover (or greening  
355 over agriculture and grazing) across northern Europe (Opdal et al. 2023).

356         Although higher riverine concentrations generally translate into higher coastal  
357 concentration, high relative differences and overall lower material concentrations in coastal sites  
358 than in their riverine counterparts indicate that most of materials are diluted and/or processed  
359 within coastal areas during summer. To better investigate the role of dilution (conservative mixing)  
360 and processing (e.g., removal, transformation, production) in coastal waters, we used mixing plots  
361 (Fig. 5) for the four coastal sectors, with each variable plotted against salinity. We observed clear  
362 differences in mixing patterns across coastal sectors, depending on the variable considered. When

363 the relationship with salinity was negative and linear ( $p < 0.05$ ), we interpreted this as a dominance  
364 of conservative mixing. In contrast, no relationship (or a positive one) suggests the influence of  
365 other processes. It is clear that phosphate behaves non-conservatively in all coastal sectors, which  
366 may explain its decoupling from river inputs. Conservative mixing of DN was more pronounced  
367 in the southern coastal sectors of the bay (C1 and C2), and seems to be the main driver of  
368  $a_{CDOM}(440)$ , the later being previously observed in the JB and Hudson Bays (Évrard et al. 2023;  
369 Meilleur et al. 2023).



371

372 **Figure 5.** Mixing plots of (A) dissolved nitrogen (DN), (B) phosphate, (C)  $\alpha_{\text{CDOM}(440)}$  and (D)  
 373 SPM in each coastal sector (represented by different colors). The lines indicate linear regressions,  
 374 with significant relationships denoted by \* ( $p < 0.05$ ) and \*\* ( $p < 0.01$ ).

375

376           Losses through flocculation and sedimentation (resulting in non-conservative mixing) of  
377 riverine organic matter and particles may also play an important role in decreasing CDOM and  
378 turbidity in JB coastal waters, as demonstrated in previous studies in the region (Stross and Sokol  
379 1989). An empirical-experimental study in Finland boreal estuaries showed that flocculation  
380 processes, induced also at low salinities, can remove up to 16% of dissolved organic carbon (DOC)  
381 (Asmala et al. 2014). Using an experimental approach, these authors demonstrated that a  
382 significant portion of DOC and dissolved iron pools are converted into particulate matter and can  
383 therefore settle through sedimentation (Asmala et al. 2014). Processing of the material can also  
384 influence properties of the dissolved organic matter (DOM), such as CDOM absorbance, and  
385 nutrients (Asmala et al. 2014, 2017).

386           Although we did not observe a clear relationship between river and coastal clusters with  
387 respect to both nutrients analyzed (N and P), potentially due to the different fractions collected  
388 (total for rivers, dissolved for the coast), we nevertheless observe a consistently 2 to 3-fold higher  
389 riverine TN concentration relative to measured coastal DN concentration for all clusters, and a  
390 lack of relationship between riverine and coastal concentrations. This might result from high rates  
391 of N<sub>2</sub> fixation by freshwater diazotrophs in rivers compared to marine waters (Blais et al. 2012),  
392 and from the rapid uptake of inorganic N by primary producers in nearshore regions (Tank et al.  
393 2012). Likewise, coastal PO<sub>4</sub> and riverine TP appeared to be largely decoupled along the eastern  
394 JB as concentrations increase along the salinity gradient following a non-conservative pattern (Fig.  
395 5D).

396           The lack of correlation between river and coastal conditions in the northern Bay sector  
397 (high relative differences for cluster 4) may be due to a mismatch in sampling locations north of  
398 the LG River due to logistical constraints. The small northernmost rivers of JB (Seal and Salmon)

399 and the small Guillaume River were sampled, but the intermediate rivers were not (Fig. 1). Coastal  
400 sampling was limited to the vicinity of the Guillaume and Piagochioui rivers (Fig. 2). Therefore,  
401 low material concentrations in coastal cluster C4 could be the result of both the dilution effect of  
402 the LG River plume and the low discharge of the Guillaume and Piagochioui rivers.

403 Coastal cluster C1 in southeastern JB receives direct and advective freshwater inputs from  
404 several large rivers. This sector of the Bay has the highest  $a_{\text{CDOM}(440)}$  and SPM, limiting light  
405 availability along the coast for primary producers such as phytoplankton, microphytobenthos,  
406 macroalgae, and the eelgrass *Zostera marina*. Despite this limitation, the surface chl *a*  
407 concentration is the highest among the different coastal sectors, with a mean value of  $2.7 \mu\text{g L}^{-1}$   
408 compared to  $1.9 \mu\text{g L}^{-1}$  in C2 and  $2.4 \mu\text{g L}^{-1}$  in C3. This indicates that the phytoplankton  
409 community is well acclimated to the low light conditions in these waters. R1 (and R0) rivers, which  
410 themselves had high chl *a* concentrations, may also contribute to the enhanced phytoplankton  
411 biomass along the coast. However, this pattern of chl *a* concentration was not the same between  
412 other coastal and river clusters, with overall higher values in coastal waters than in rivers (on  
413 average ~9% higher). This suggests that rivers enrich coastal waters with dissolved inorganic and  
414 organic nutrients and promote the growth of coastal primary producers, rather than simply  
415 exporting chl *a*.

416 We acknowledge several limitations of this study. First, samples were not taken  
417 simultaneously in rivers and coastal sites, which could introduce temporal decoupling in  
418 physicochemical conditions and noise in the relationship between rivers and nearshore  
419 environments. To mitigate this issue, we averaged data from two consecutive summers for rivers  
420 and included data from multiple sites. We recognize that processes such as shoreline erosion and  
421 sediment resuspension, particularly by storm waves or fast streamflow associated with

422 hydroelectric production in the LG River, could contribute additional SPM and nutrients to the  
423 coastal waters. It is important not to disregard the role of LG as a sporadic source of sediments to  
424 the coastal system due to high erosion of downstream riverbanks close to the mouth during high  
425 flow events (Lefebvre et al. 1991). This could explain the few observations where mean SPM or  
426 phosphorus in coastal clusters were higher in coastal waters than in the paired river cluster (cluster  
427 3 and 2, respectively). Secondly, this study only covered the summer season, because there was  
428 no matching oceanographic data for spring and fall due to logistic impediments in sampling the  
429 Bay waters, and we acknowledge that seasonality can play an important role in determining  
430 riverine influence on the coast. During winter, the typically low discharge or complete freezing of  
431 unregulated rivers may lead to reduced riverine influence along the coast, whereas in spring,  
432 increased river discharge due to snowmelt likely enhances the influence of rivers on coastal waters.  
433 A recent study demonstrated that nitrate and phosphate distributions in the La Grande plume  
434 (extended north and to a lesser extent south of the LG mouth) is driven by conservative mixing in  
435 winter, while that in summer, both water mass mixing and biological nutrient uptake contribute to  
436 their availability (Guzzi et al. 2024).

437         In conclusion, our study identified significant latitudinal variability in the physicochemical  
438 properties of both rivers and coastal sites along the eastern James Bay. While the volume of river  
439 water discharged into the bay plays a key role—where greater discharge leads to a stronger riverine  
440 influence on coastal waters—it is not only discharge that drives this variability. The distinct  
441 physicochemical differences between watersheds along the latitudinal gradient also play a crucial  
442 role, resulting in well-defined transition zones between river and coastal waters.

443         Given the complex interactions between terrestrial and marine systems, this study presents  
444 an integrated approach to identify transition zones and assess the impact of river exports on

445 physicochemical conditions of coastal waters along the eastern JB, which can be adapted to river-  
446 coast continua elsewhere. Given the anticipated influence of climate and environmental changes  
447 on river exports globally, forthcoming research should focus on assessing the effects of short- and  
448 long-term disturbances on coastal health, resilience, and the key ecosystem services provided by  
449 these environments.

450

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465

466 **Author Contribution Statement**

467 PdG, UN and MG acquired the funds. PdG, MLG and CFM designed the field survey for rivers,  
468 while UN, MG, VG and HX designed the field survey along the coast. CFM, VG, and MLG along  
469 other contributors conducted field work. MdM and CFM designed the study approach, analysed  
470 the data, and conducted the statistical analyses. MdM wrote the first draft, and all co-authors  
471 discussed and edited the subsequent versions.

#### 472 **Conflict of Interest**

473 Authors declare no conflict of interest.

#### 474 **Data Availability Statement**

475 Metadata will be published in an open access repository (Borealis platform).

476 Additional Supporting Information may be found in the online version of this article.

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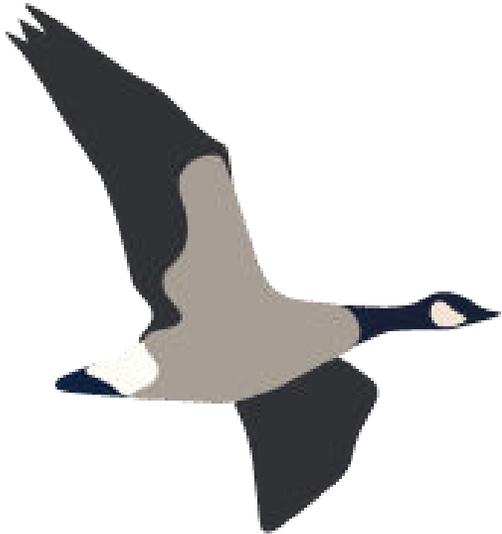
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616

Steering Committee

January 29 2025

**PRESENTING CHCRP PHASE II (2023-2024)  
FINAL REPORT**



Mary O'Connor, Fanny Noisette, Zou Zou Kuzyk, Jens Ehn, Simon Bélanger, Caroline Fink-Mercier, Nicole Knight

# Timeline of the different phases of the CHCRP



Building Alliance project in response to community priorities

# Timeline of the different phases of the CHCRP

Continue monitoring efforts started in Phase I  
and prepare the Alliance proposal



Building Alliance project in response to community priorities

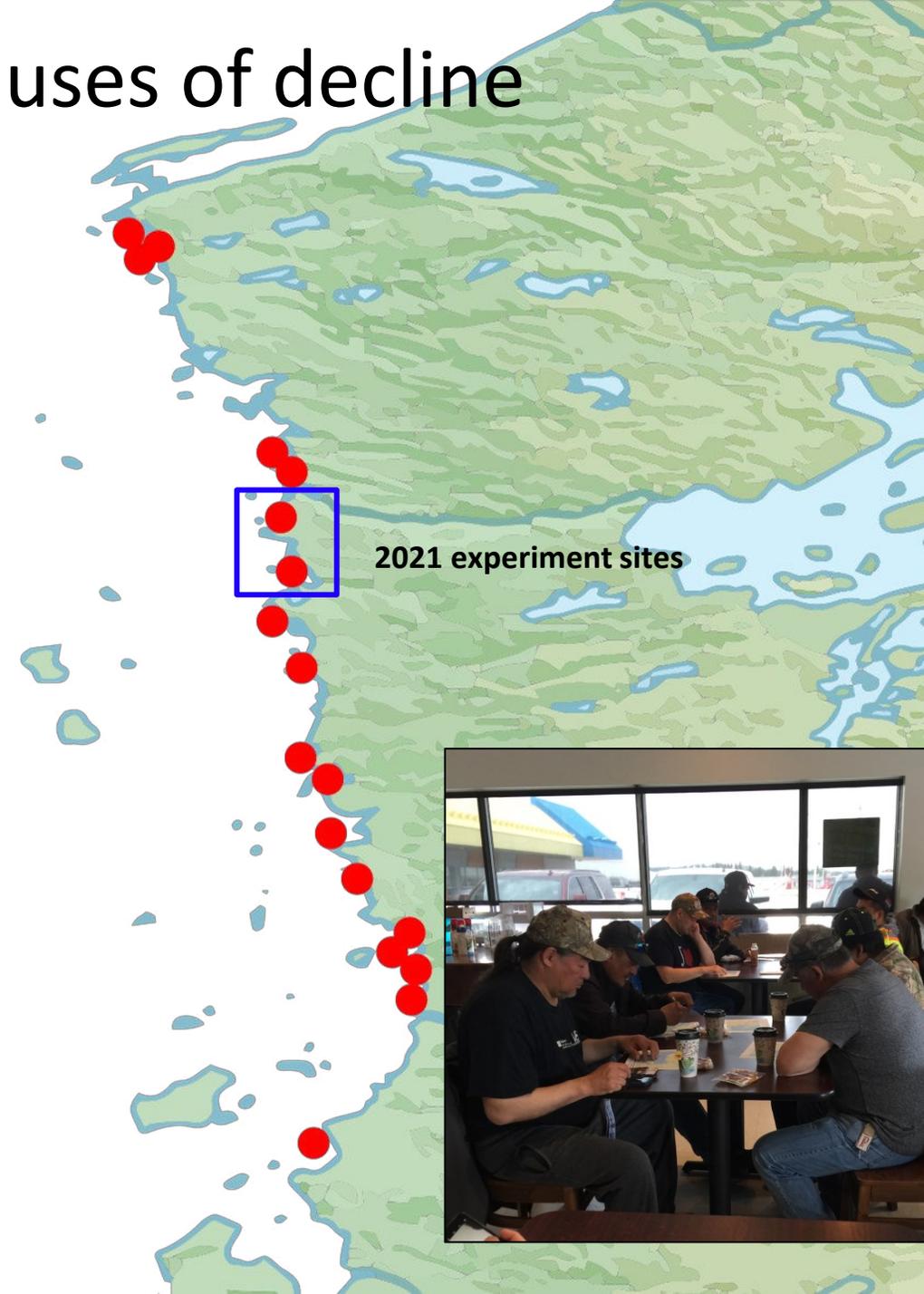
# Timeline of the different phases of the CHCRP



Final Eelgrass Report  
presented previously by  
Mary O'Connor

# Eelgrass condition, causes of decline

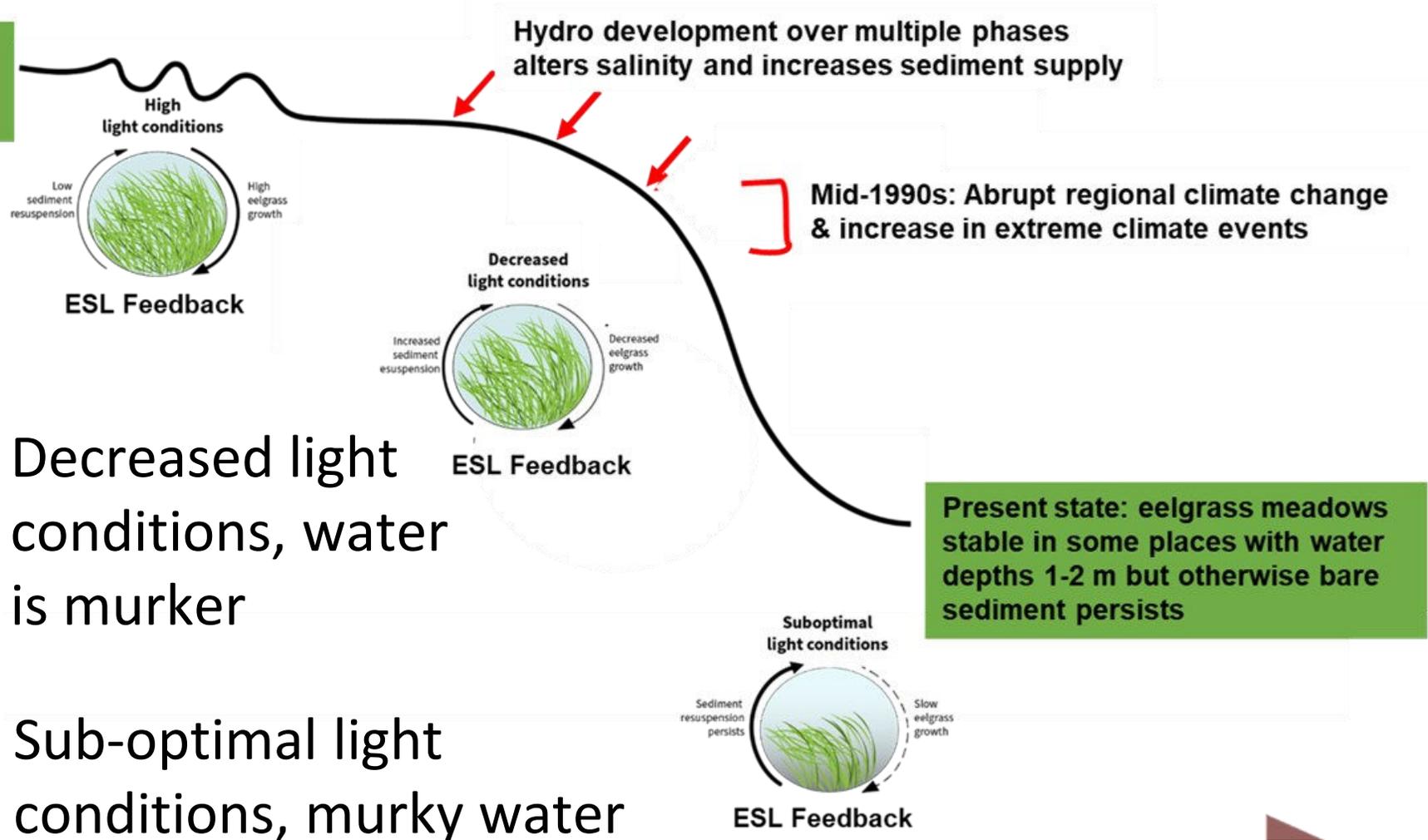
- 124 sites (2019, 2020, 2021) - eelgrass team
- > 700 sites (2017-2021) - ocean team
- We did experiments to test for effects of light and nutrient changes
- We measured eelgrass and biodiversity



# Eelgrass persists but is not recovered

Initial state: eelgrass meadows widespread and stable in water depths 1-4 m

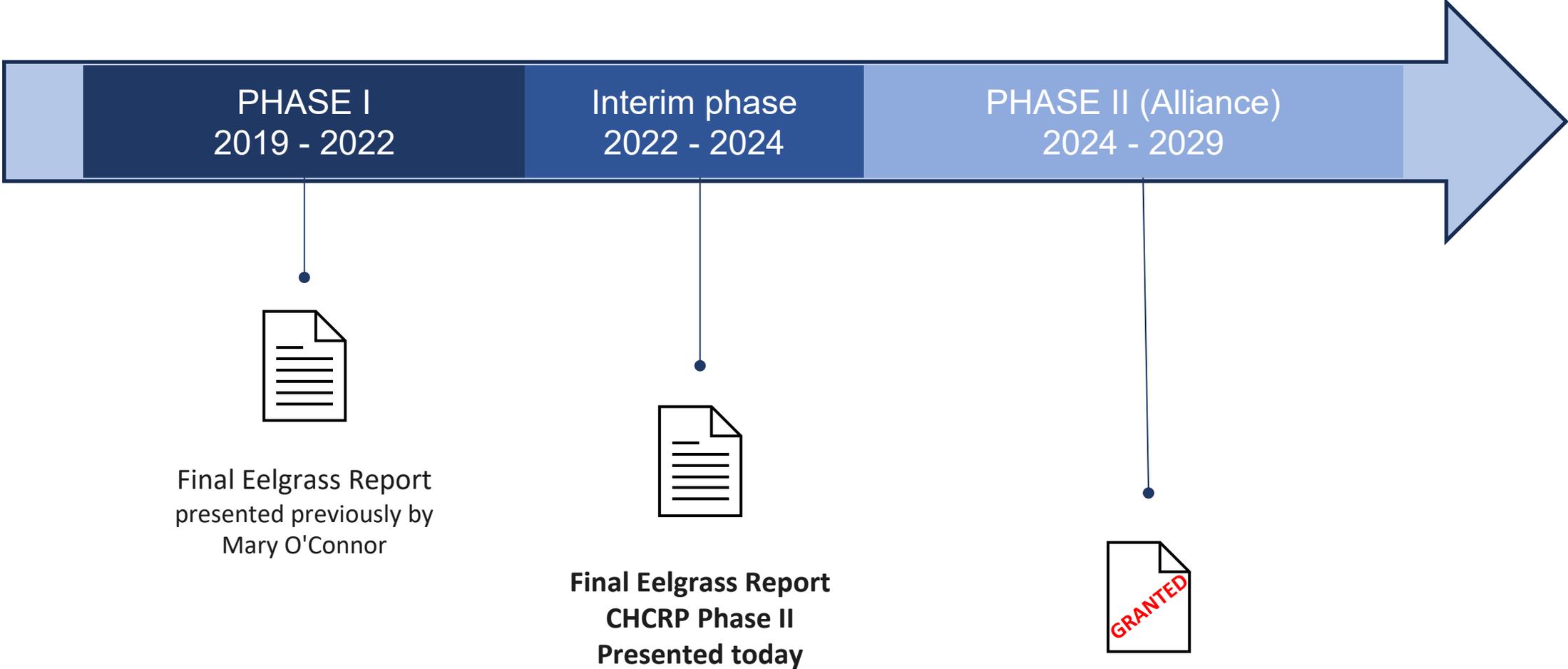
High light conditions before 1990; water was clearer



Decreased light conditions, water is murker

Sub-optimal light conditions, murky water

# Timeline of the different phases of the CHCRP

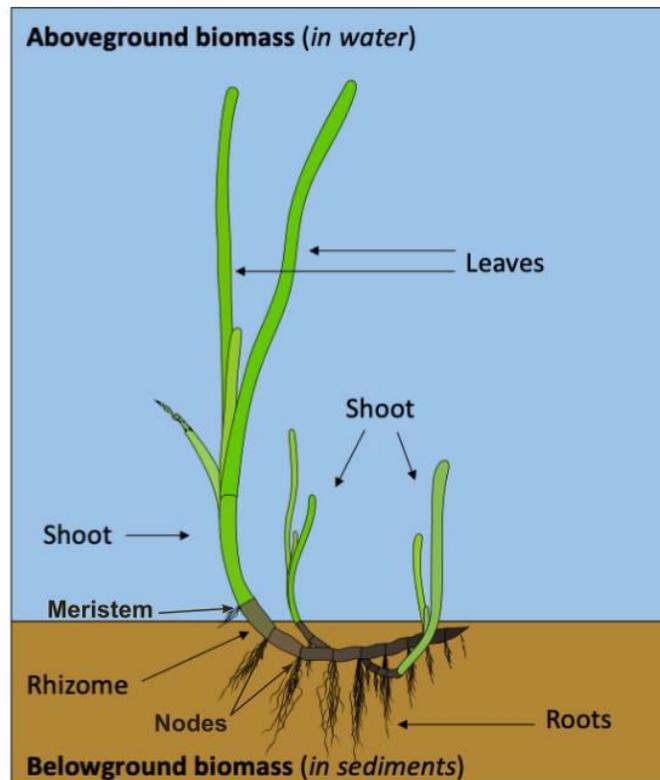


# Objectives

- **Obj 1:** To pilot field and laboratory methods that will be used to address Phase II goals
- **Obj 2:** To assess eelgrass condition at key locations by visiting monitoring sites that have been visited several times during the previous years and/or that were of particular interest to land users
- **Obj 3:** To follow eelgrass growth at 5 locations to understand the link between early summer growth and environmental conditions
- **Obj 4:** To co-develop overall goals and objectives of Phase II with land users, university partners and community leadership, and write and submit a NSERC/CRSNG proposal for funding.

# Summer 2023

- Fieldwork in Eeyou Istchee cancelled because of forest fires
- Pilot methods useful for phase 3: Rhizome growth and sugar contents (obj 1)

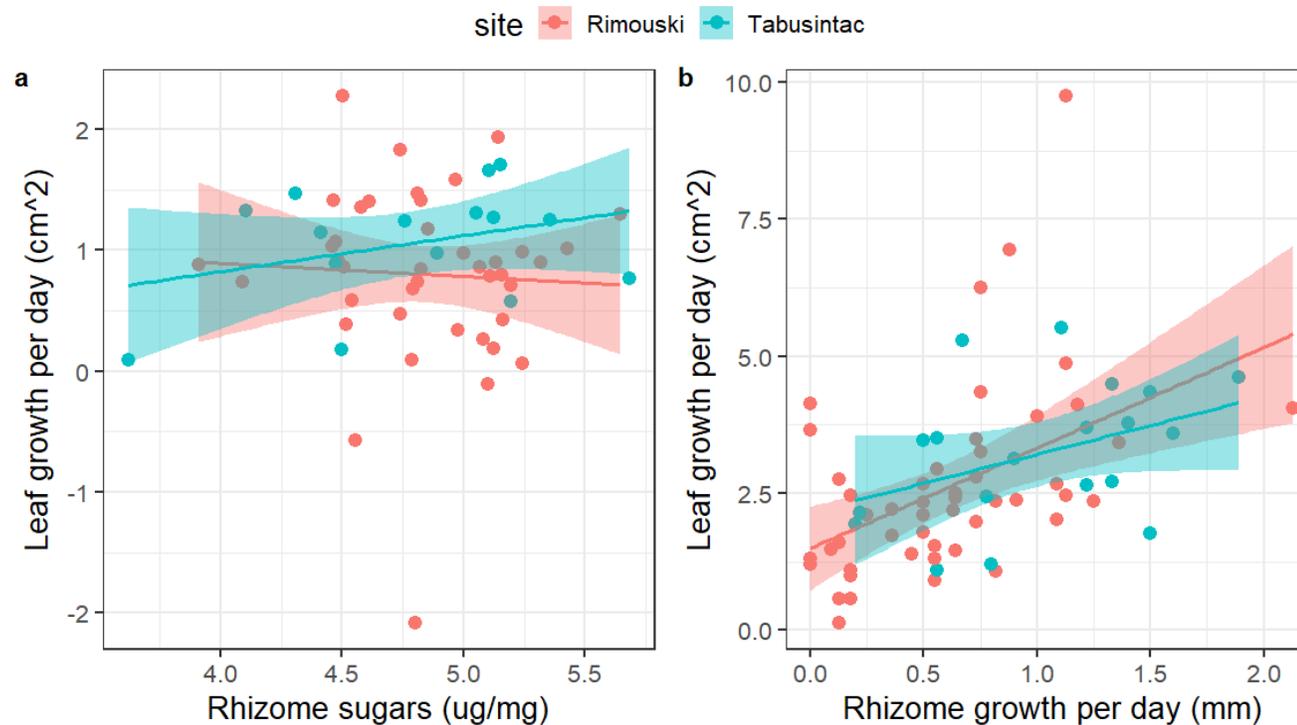


Rhizome growth: tool to understand the influences of environment on eelgrass growth and survival even during time periods when sampling is difficult (e.g. winter)

Rhizome sugar content: to quantify storage that can be used to overwinter and start growth in early spring

# Summer 2023

- Fieldwork in Eeyou Istchee cancelled because of forest fires
- Pilot methods useful for phase 3: Rhizome growth and sugar contents (obj 1)



Interim phase  
2022 - 2024

# Summer 2024



# Summer 2024

- Fieldwork in Eeyou Istchee (June 24-July 22 )
- Community meetings to discuss project and priorities



# Summer 2024

- To assess eelgrass condition at key locations (obj 2)

## Seagrass measurements

Presence/Absence

Leaf morphometry

Biomass

Rhizome morphometry

Leaf physiology

Density

% Cover

## Water column measurements

Nutrients

CDOM

Chla

SPM (suspended particulate matter)

## Sediment measurements

Pore water nutrients

Sediment hardness

Sediment core (for grain size)

Temperature, salinity, pH, Oxygen

## Trapline

## Participating tallymen and land users

CH34

Louis House, Charles House, Lawrence Napash, Darren Stephen

CH33

John Sam

CH03

John Rupert, Ronnie Rupert, Lawrence Napash

VC10

Rene Atsynia, Henry Stewart, Leonard Asquabaneskum

VC11

Rene Atsynia, Henry Stewart, Leonard Asquabaneskum, Roland Tomatuk

VC12

Roland Tomatuk, Abraham Matches, Cody Mark, Rene Atsynia

VC17

Ernie Hughboy, Stanley Shashaweskum

VC14

Henry Stewart, Leonard Asquabaneskum

VC13

Henry Stewart, Leonard Asquabaneskum

VC32

Marcel Moses, Wilfred Cheezo

VC15

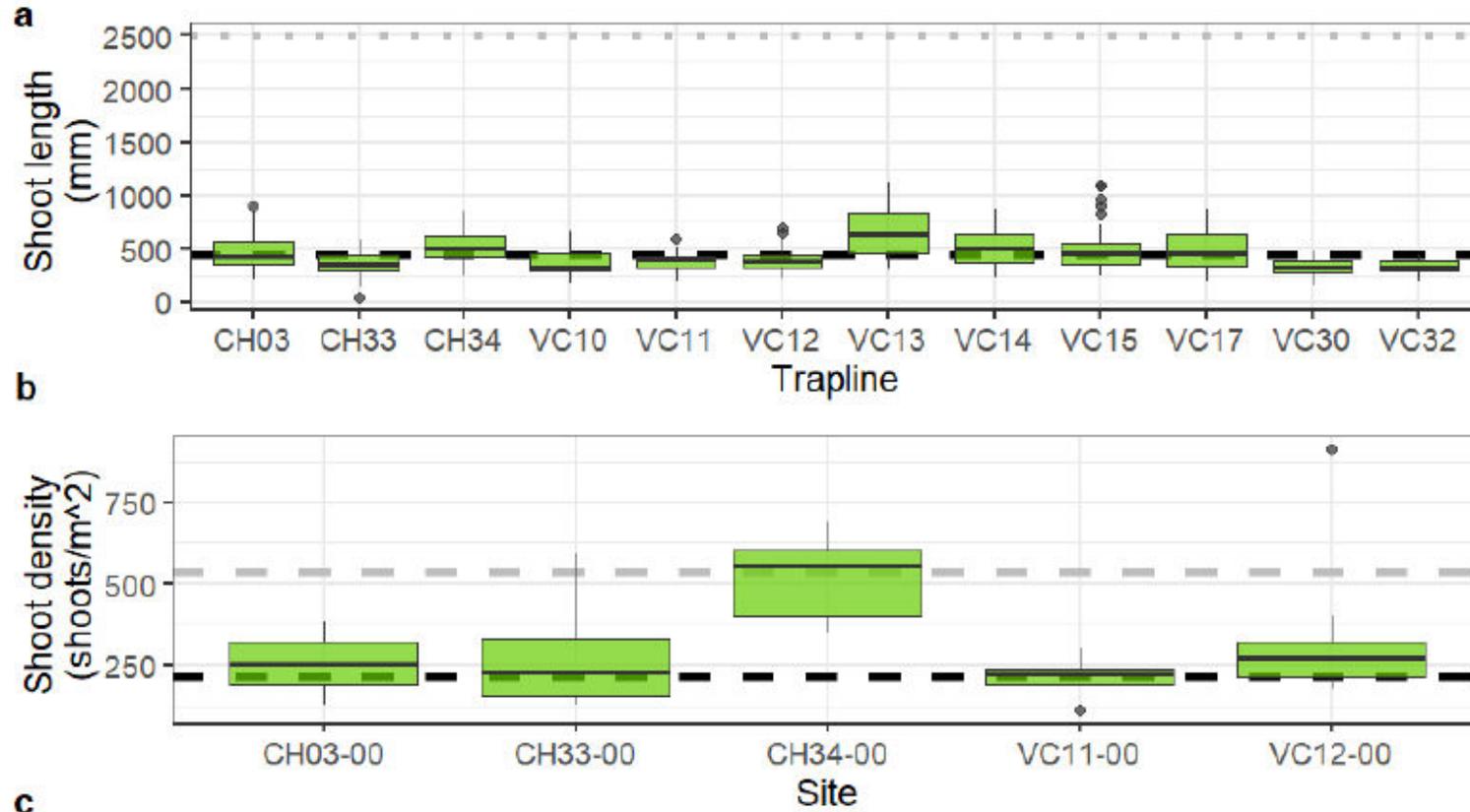
Marcel Moses, Wilfred Cheezo

VC30

Marcel Moses, Wilfred Cheezo

# Summer 2024

- To assess eelgrass condition at key locations (obj 2)

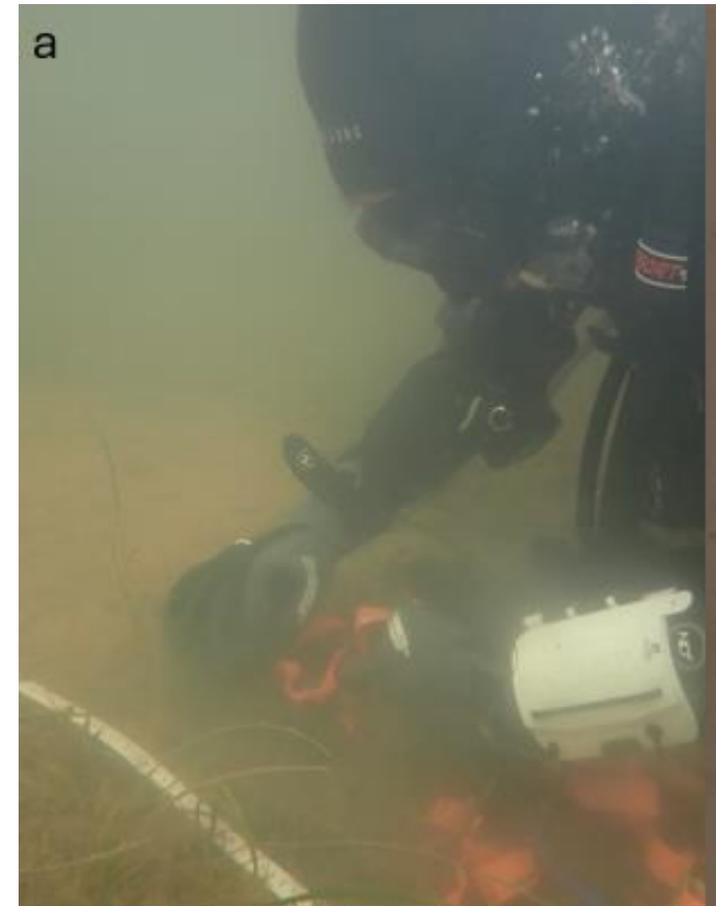
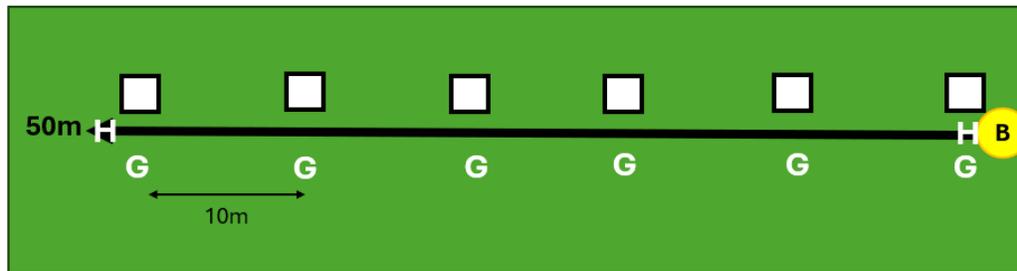


Eelgrass shoot lengths, density, and biomass were all comparable to regional averages in recent years but generally remain below pre-decline regional averages

# Summer 2024

- To follow eelgrass growth at 5 locations to understand the link between early summer growth and environmental conditions (obj 3)

- B** Surface buoy
- Density, biomass, sediment, porewater
- G** Growth
- H** HOBO (measures light and temperature during the growing period)



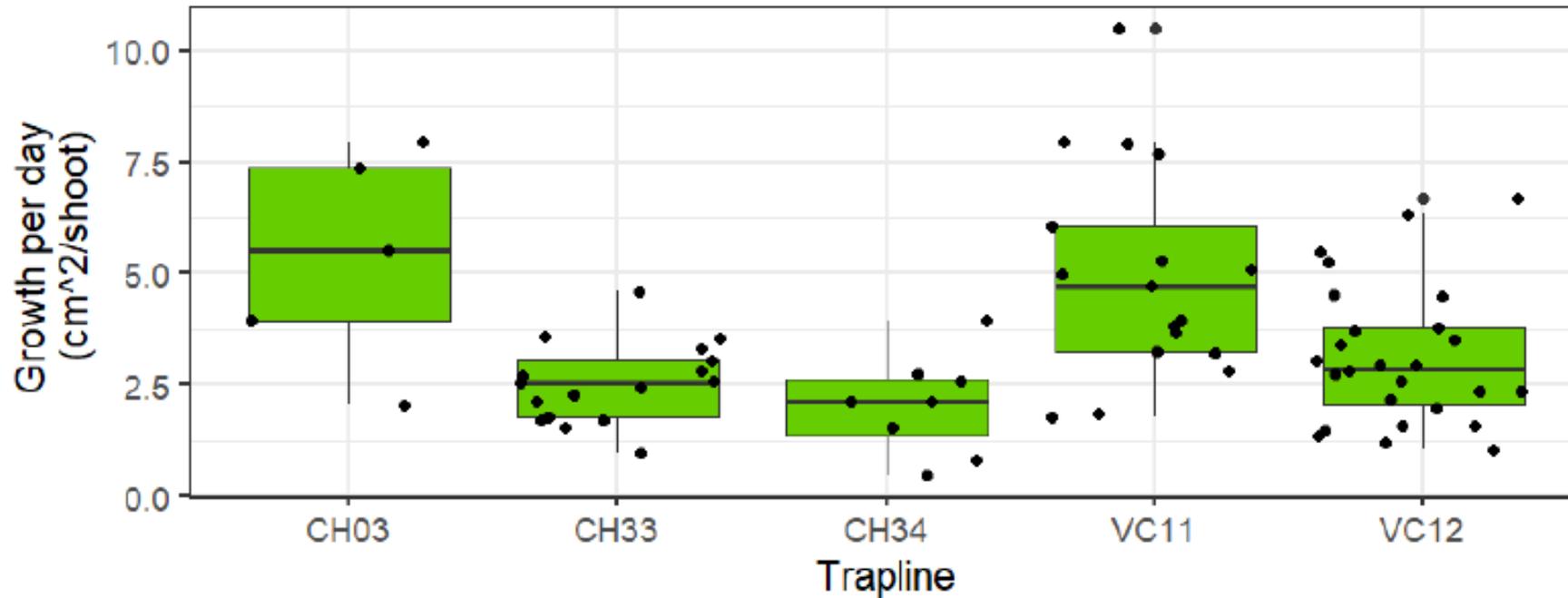
# Summer 2024

- To follow eelgrass growth at 5 locations to understand the link between early summer growth and environmental conditions (obj 3)



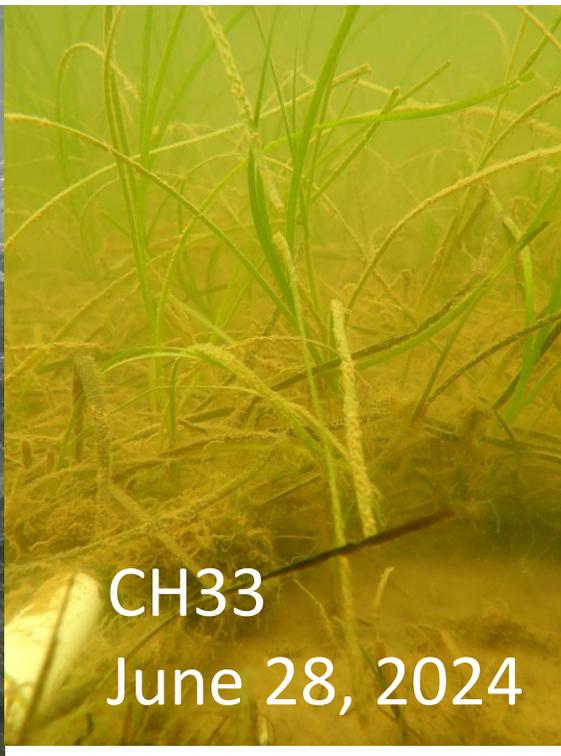
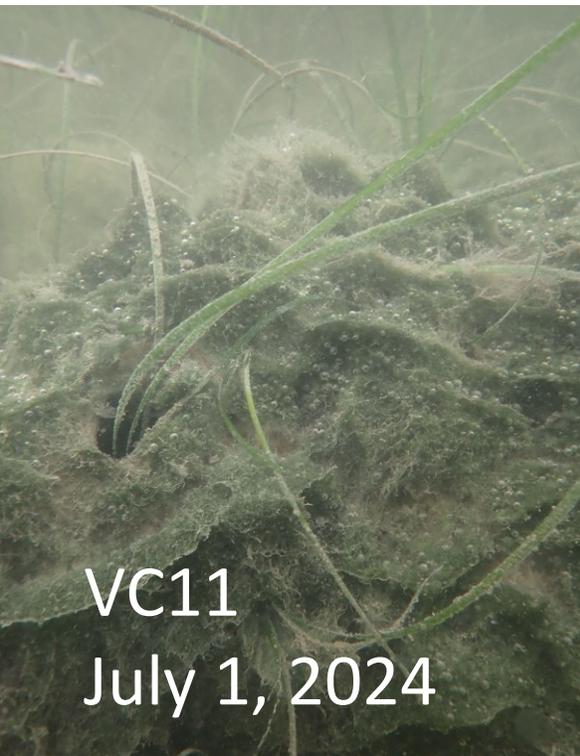
# Summer 2024

- To follow eelgrass growth at 5 locations to understand the link between early summer growth and environmental conditions (obj 3)



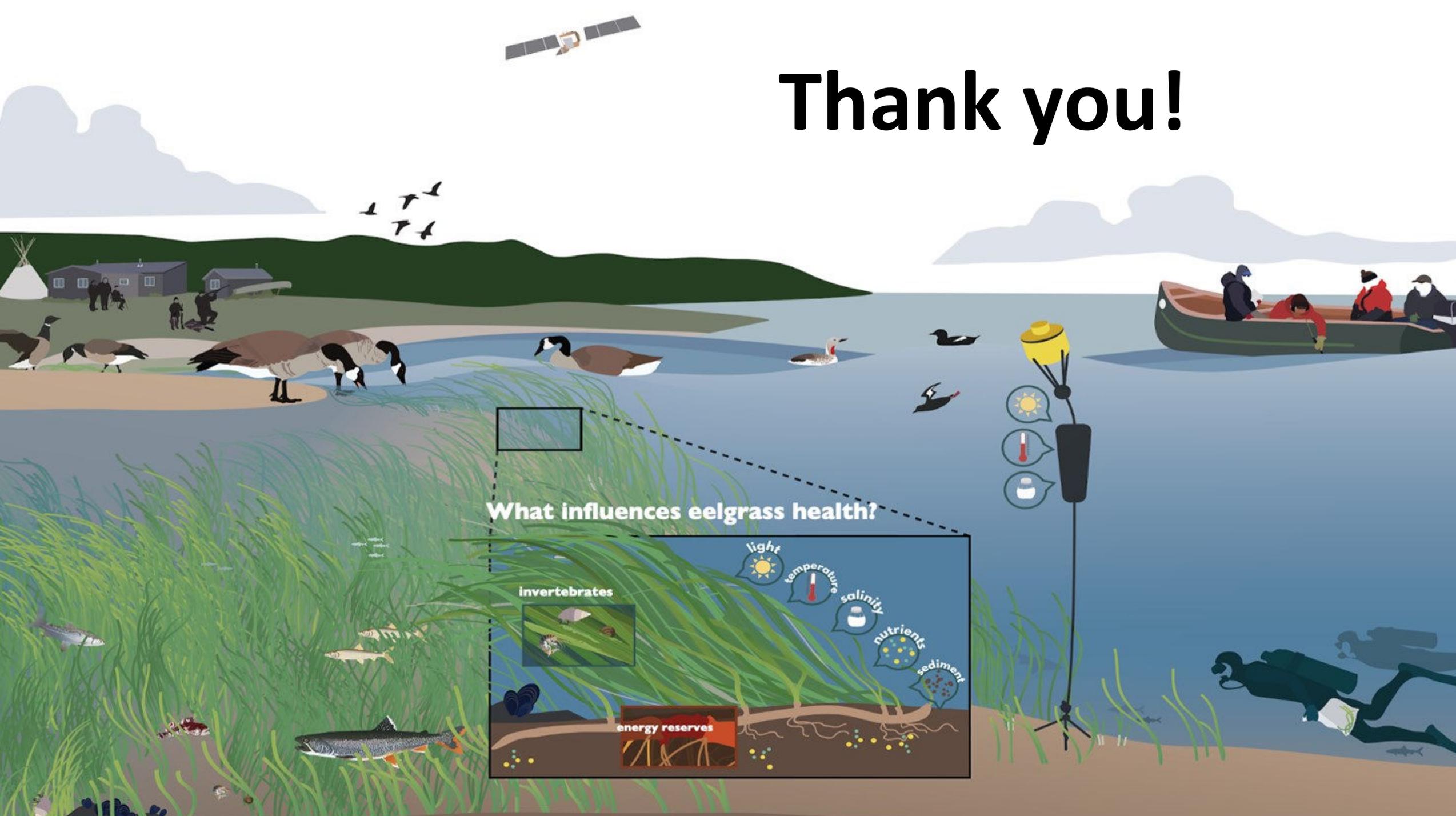
# Summer 2024

- We have seen many algal mats that may be affecting eelgrass growth

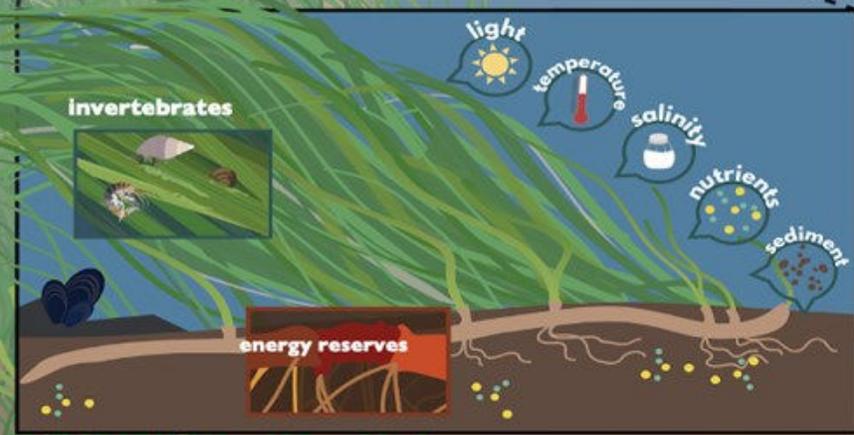




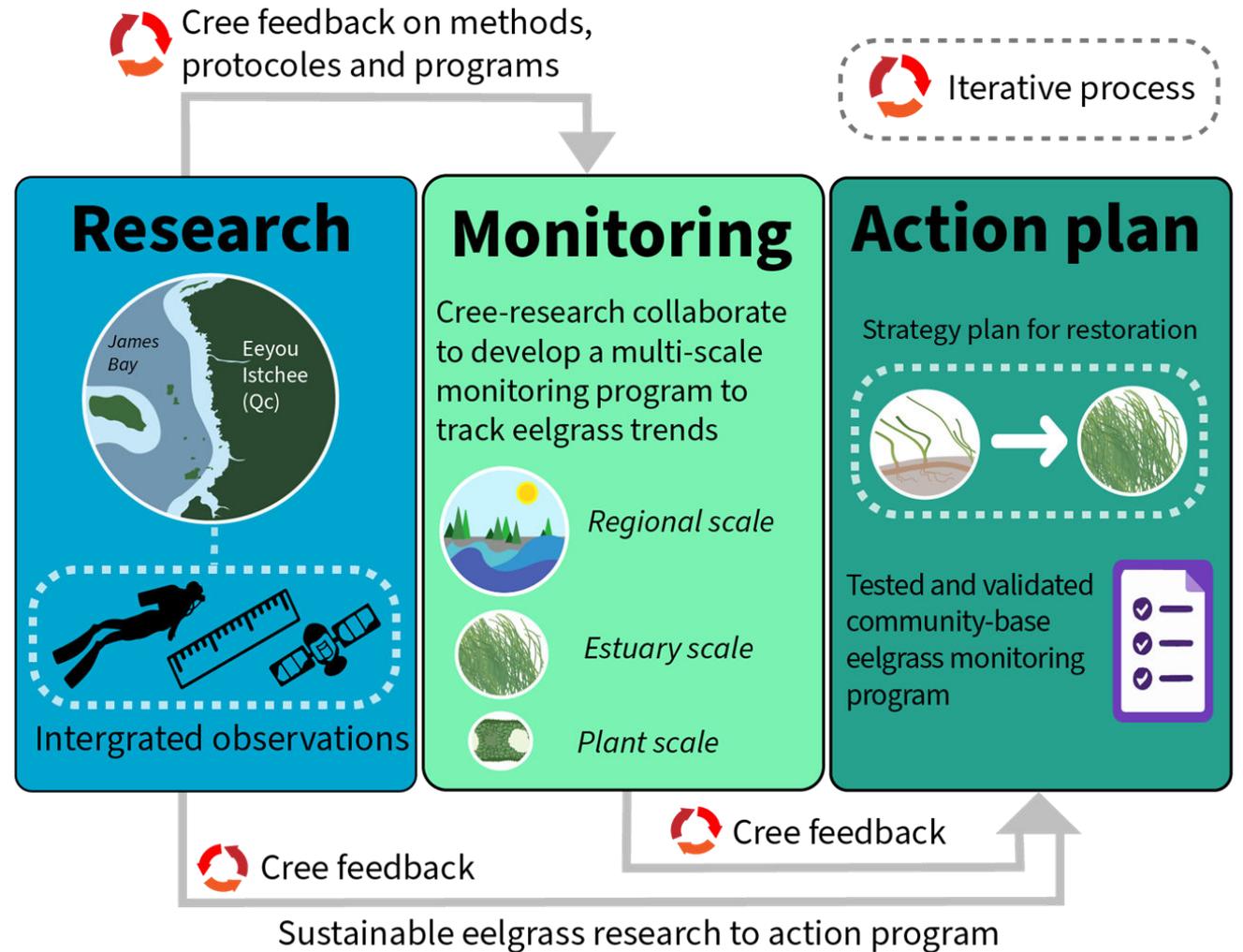
# Thank you!



What influences eelgrass health?



# Coastal Ecosystems Project Goals



# Coastal Ecosystems Project Team



## Collaborators:

Murray Humphries,  
Ally Menzies,  
Paul del Giorgio

Mary O'Connor



Jens Ehn



## co-PI

Zou Zou Kuzyk



Fanny Noisette



Simon Bélanger



## Partners (current)

Melanie Leblanc (Niskamoon)

Eeyou Marine Region Wildlife Board  
Parks Canada  
CEGRIM  
Hydro Québec

## Project coordinator Post-doc fellow

Caroline Fink-Mercier Nicole Knight



# Coastal Ecosystems Project: Goals

- 1: To understand how the ocean, river plume, and ice environment affects eelgrass health (current and future).**
- 2. To identify opportunities to enhance recovery and restoration.**
- 3: To support the development of a sustainable Cree-led eelgrass ecosystem monitoring and early action program**

